



Intergovernmental
Oceanographic
Commission



UNESCO-IOC / NOAA ITIC Training Program in Hawaii (ITP-TEWS Hawaii)

TSUNAMI EARLY WARNING SYSTEMS

AND THE PACIFIC TSUNAMI WARNING CENTER (PTWC) ENHANCED PRODUCTS

TSUNAMI EVACUATION PLANNING AND UNESCO IOC TSUNAMI READY PROGRAMME

15-26 September 2025, Honolulu, Hawaii

Tsunami Warning Center Operations: Real-Time Earthquake Detection and Fast Source Characterization – Locating Earthquakes & Estimating Magnitude

Laura Kong

UNESCO/IOC – NOAA International Tsunami Information Center

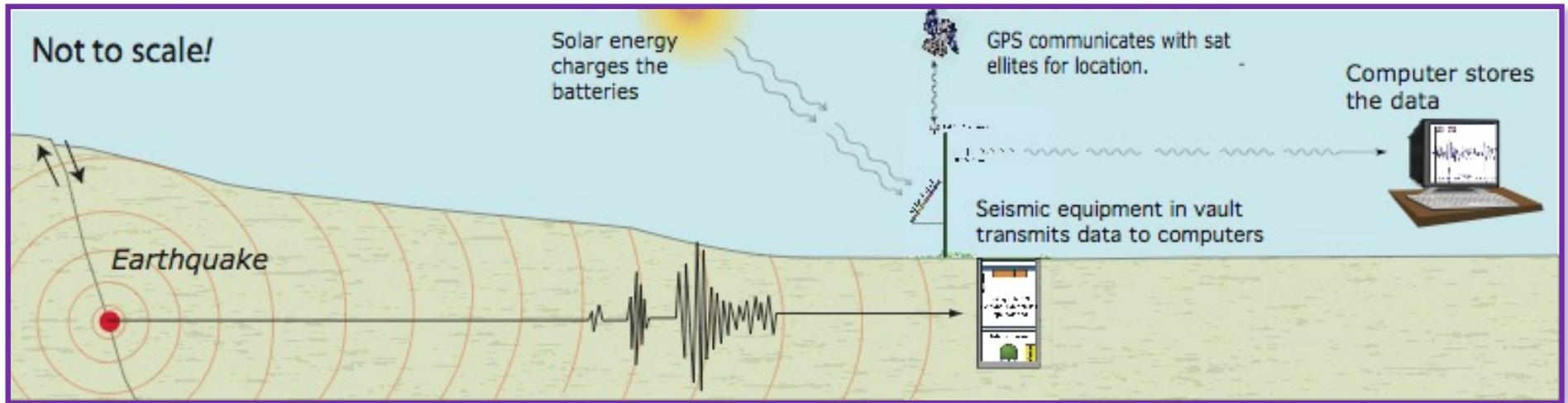
Kanoa Koyanagi, Jonathan Weiss, Stuart Weinstein, Nathan Becker
and colleagues
NOAA Pacific Tsunami Warning Center



Pacific
Community
Communauté
du Pacifique

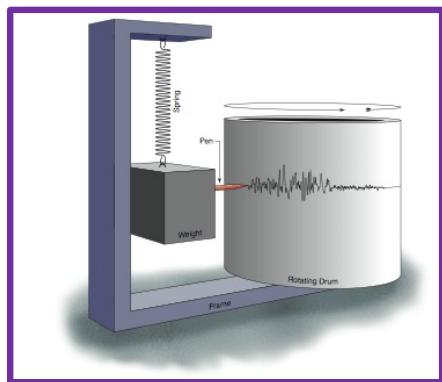
Locating Earthquakes

How do we measure earthquakes?

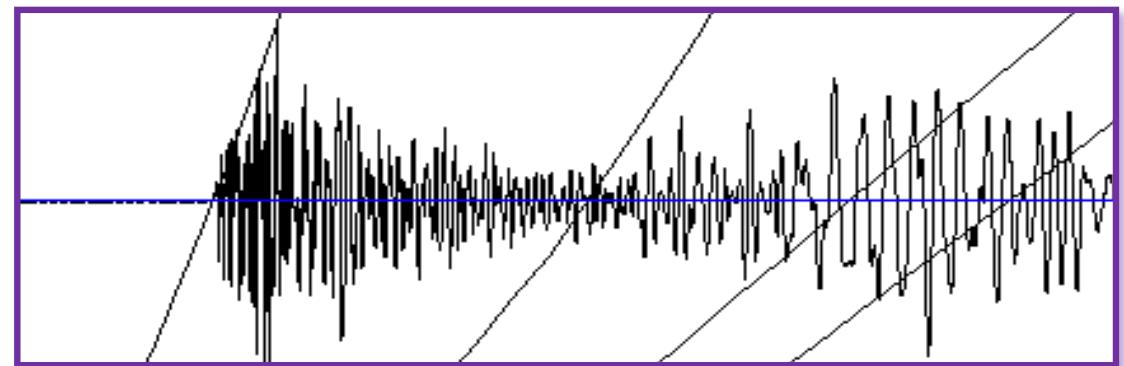


When an earthquake occurs, the seismic waves travel through the Earth to the seismic station where the information is transmitted to distant computers.

A **seismograph** detects and records earthquakes



A **seismogram** is the earthquake record



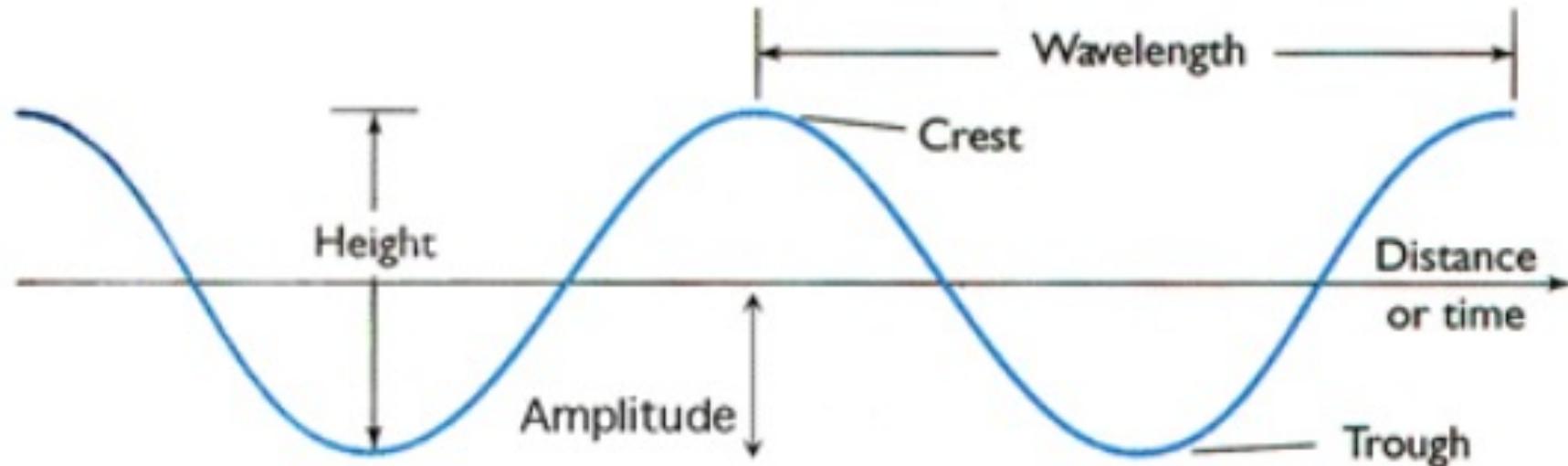
Global Seismic Network



PACIFIC TSUNAMI
WARNING CENTER

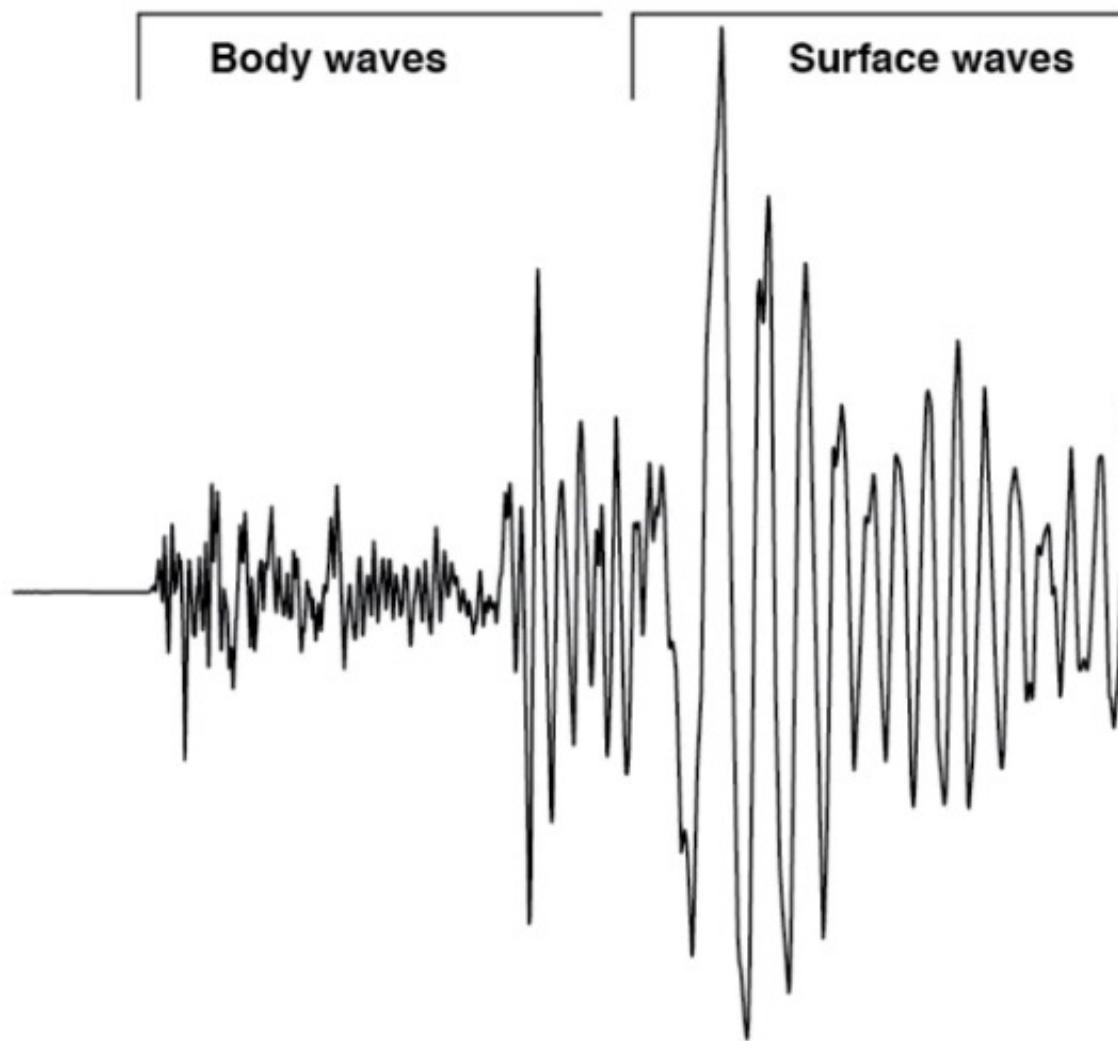
PTWC monitors >800 seismic stations worldwide

Basic (Seismic) Wave Properties



- Wavelength - distance between successive crests of the wave
- Wave period - time it takes for the wave to travel one wavelength

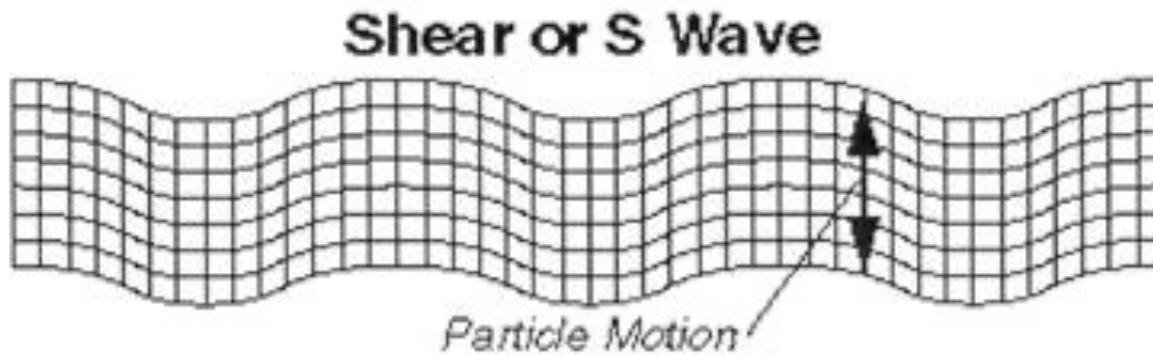
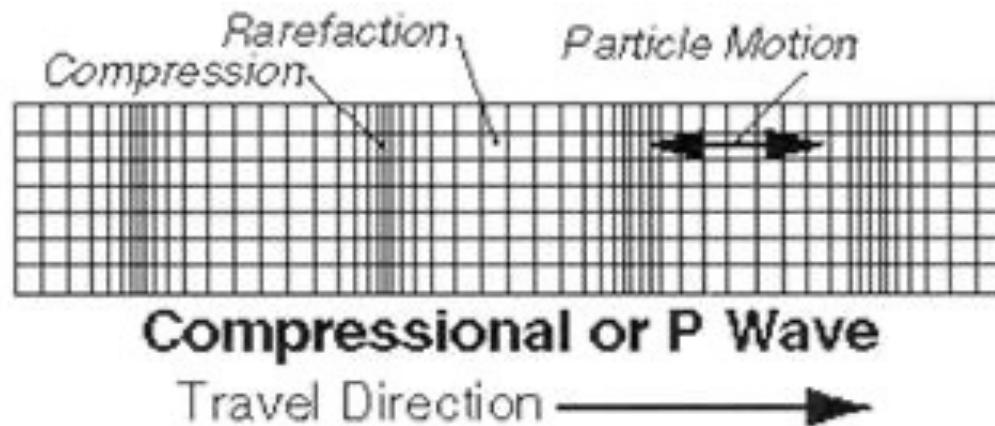
Basic Types of Seismic Waves



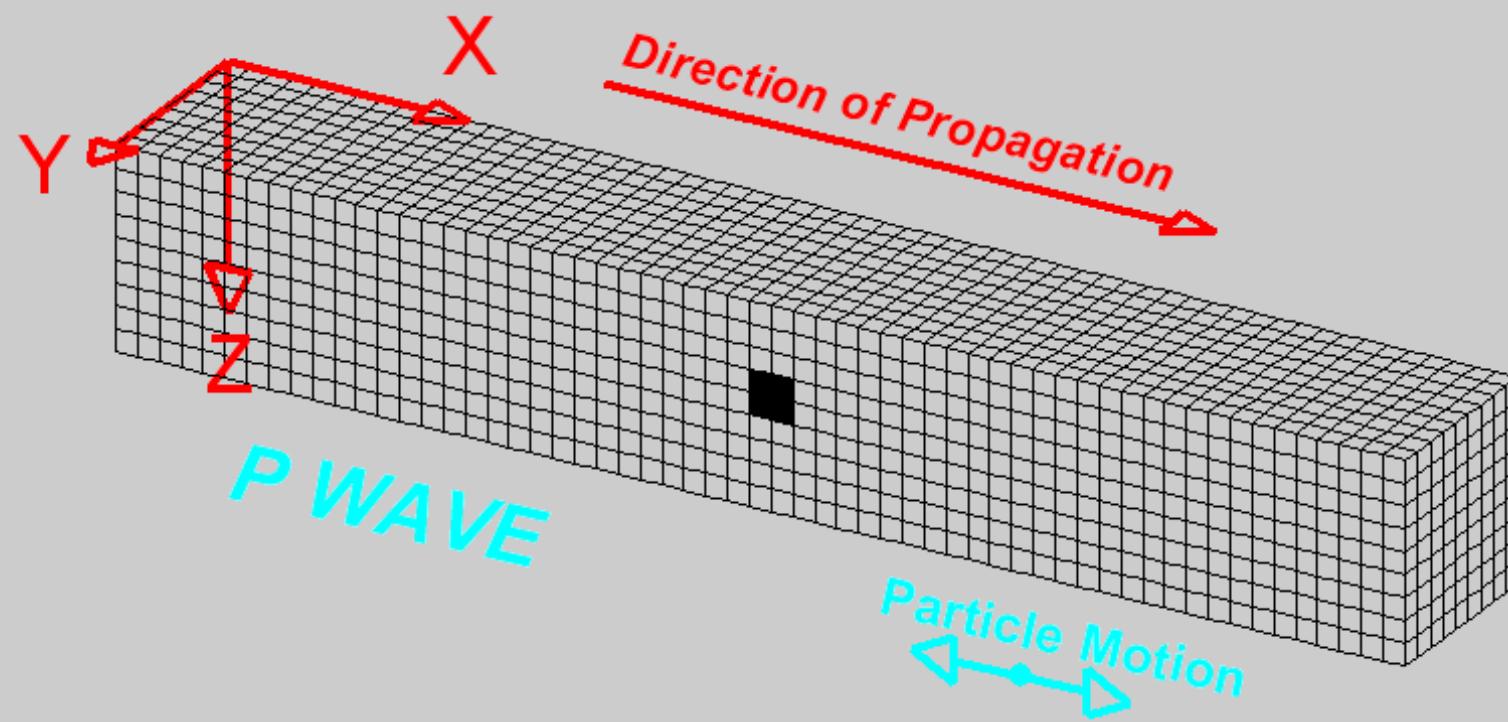
Surface waves arrive after the body waves. They have lower frequency and larger amplitude than body waves.

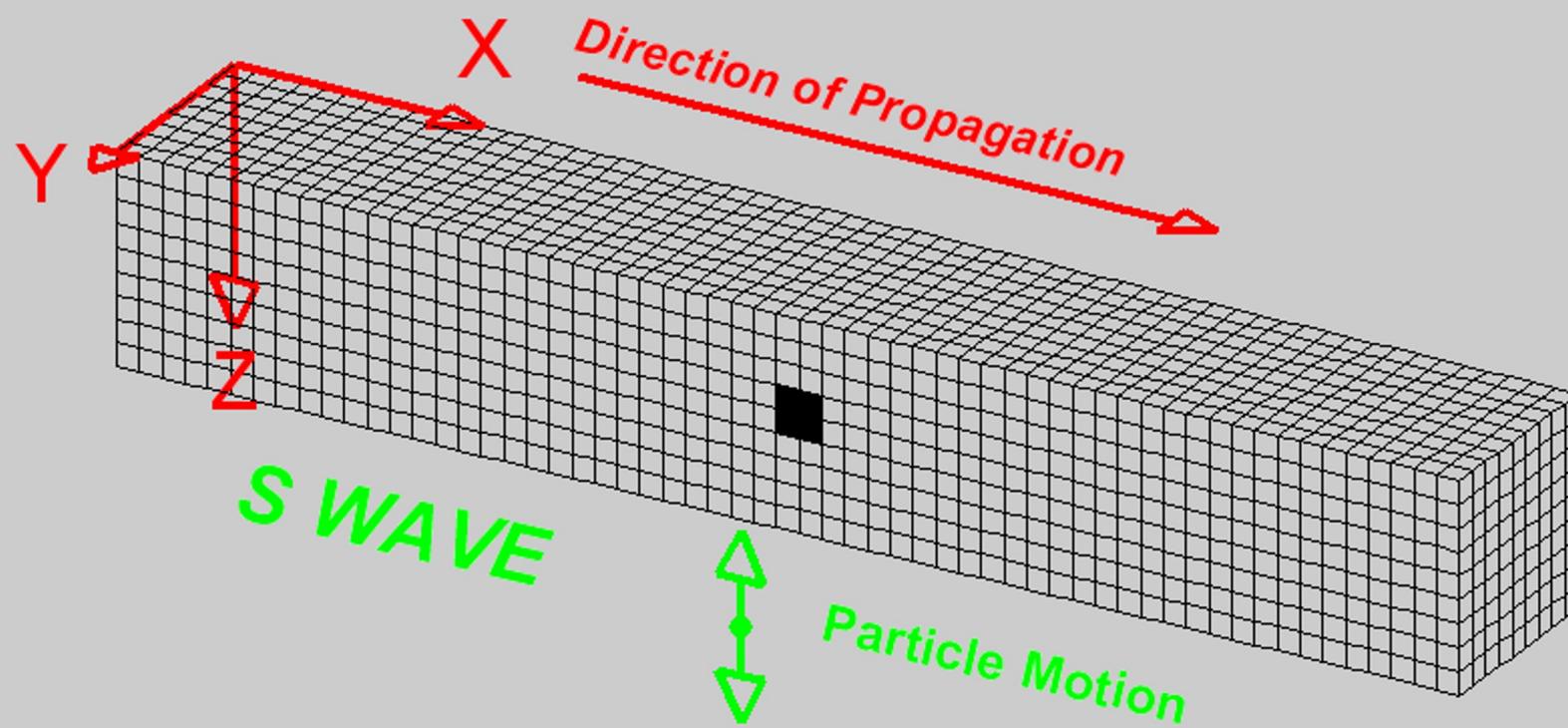
Basic Types of Seismic Waves

P and S Waves (also called body waves)



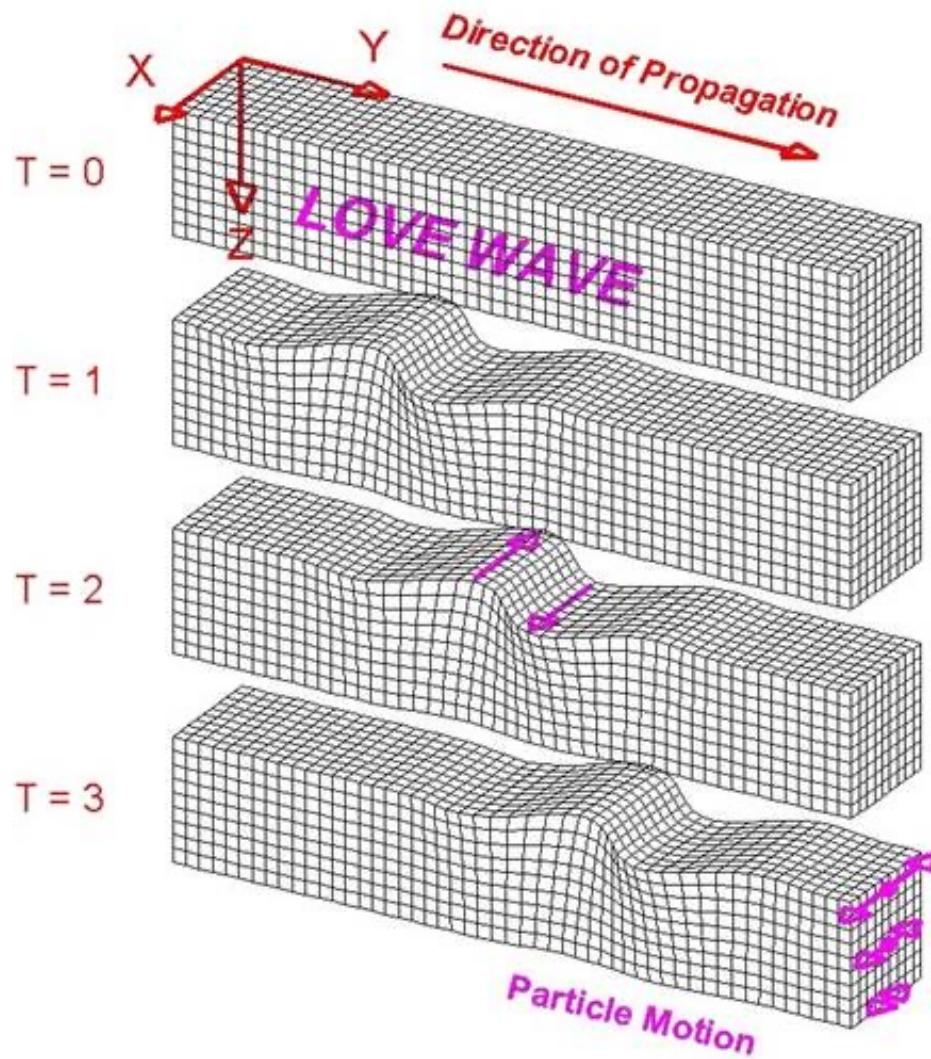
P waves are faster and travel at speeds of **6-14 km/s**



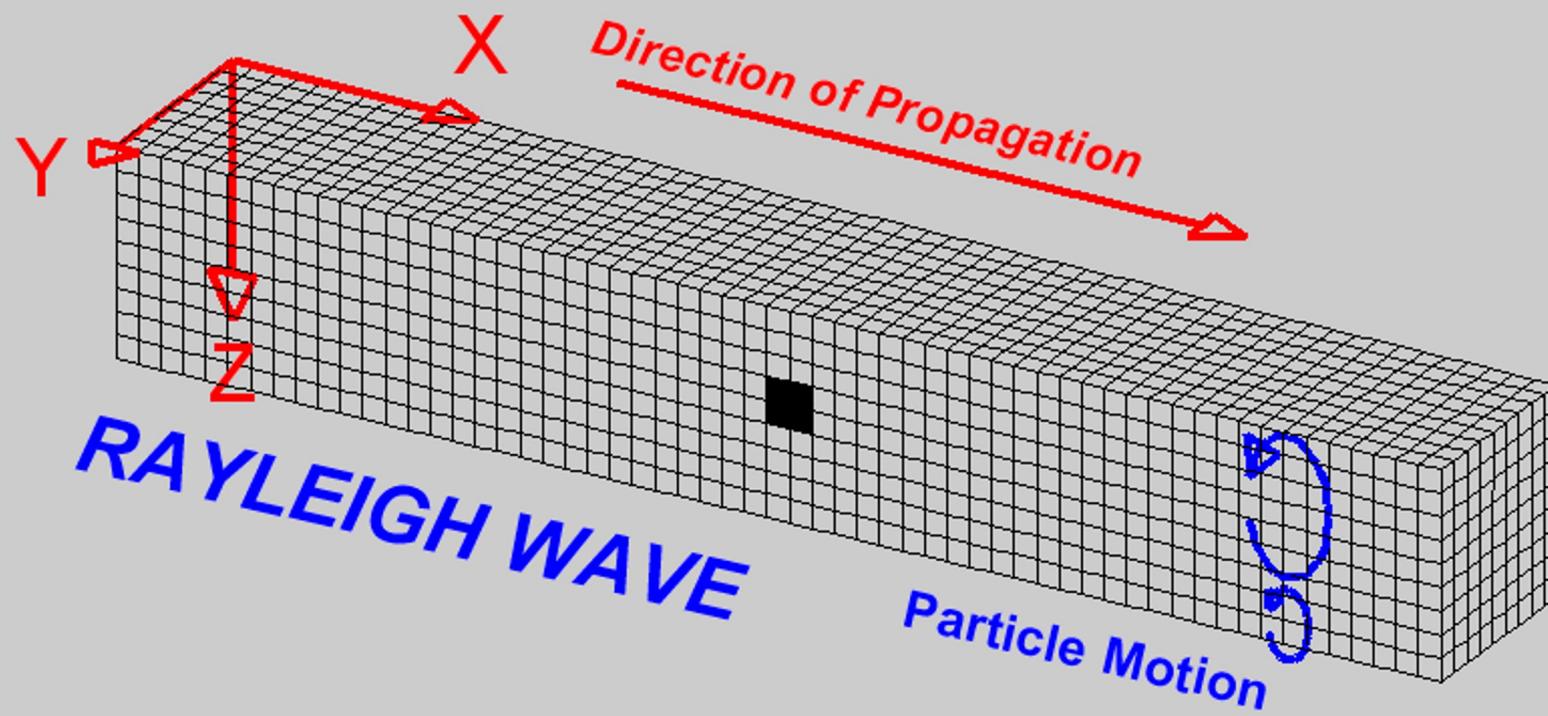


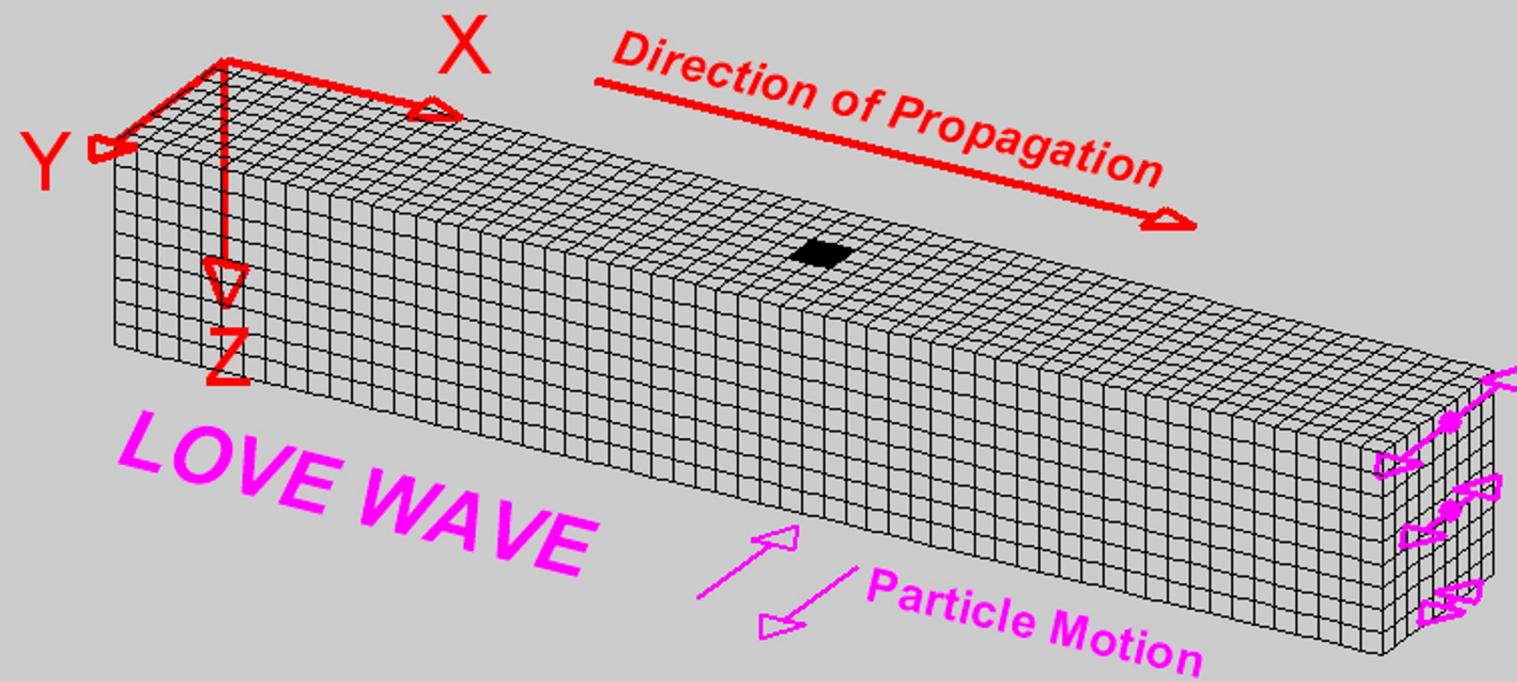
Basic Types of Seismic Waves

Surface Waves (do not travel through the earth)

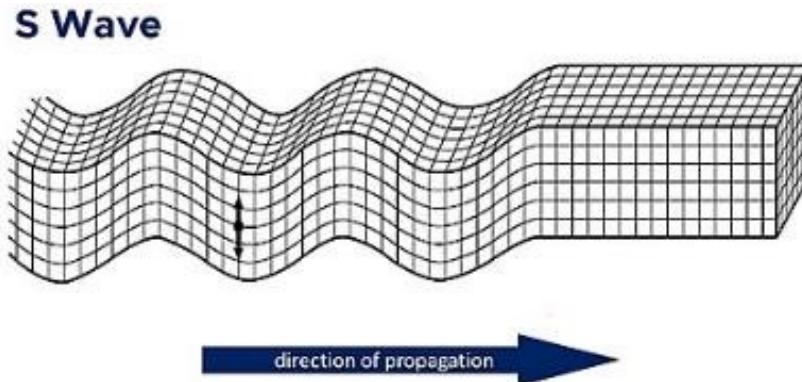
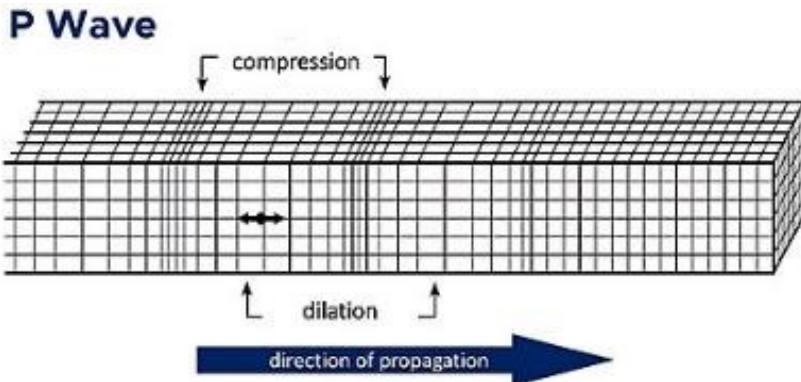


Surface waves travel along the Earth's surface at speeds of $\sim 3/\text{km/sec}$, much slower than the body waves.

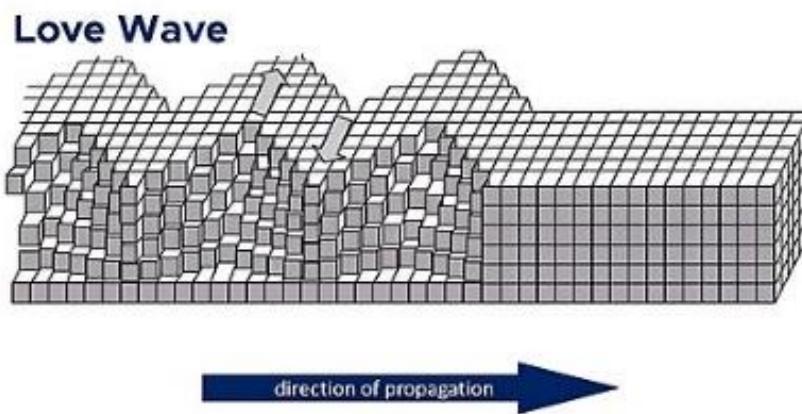
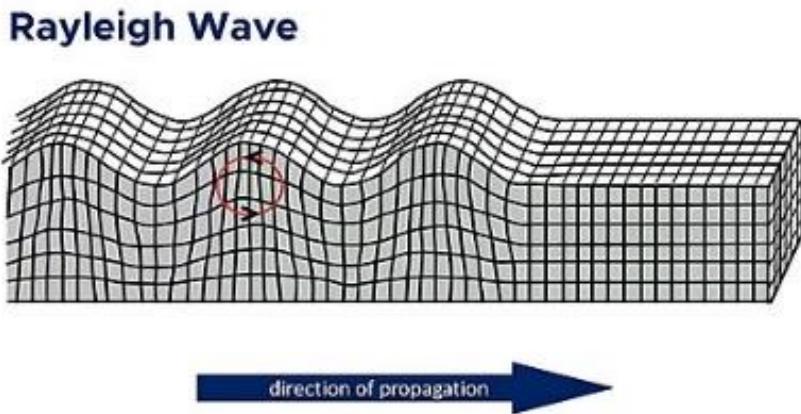




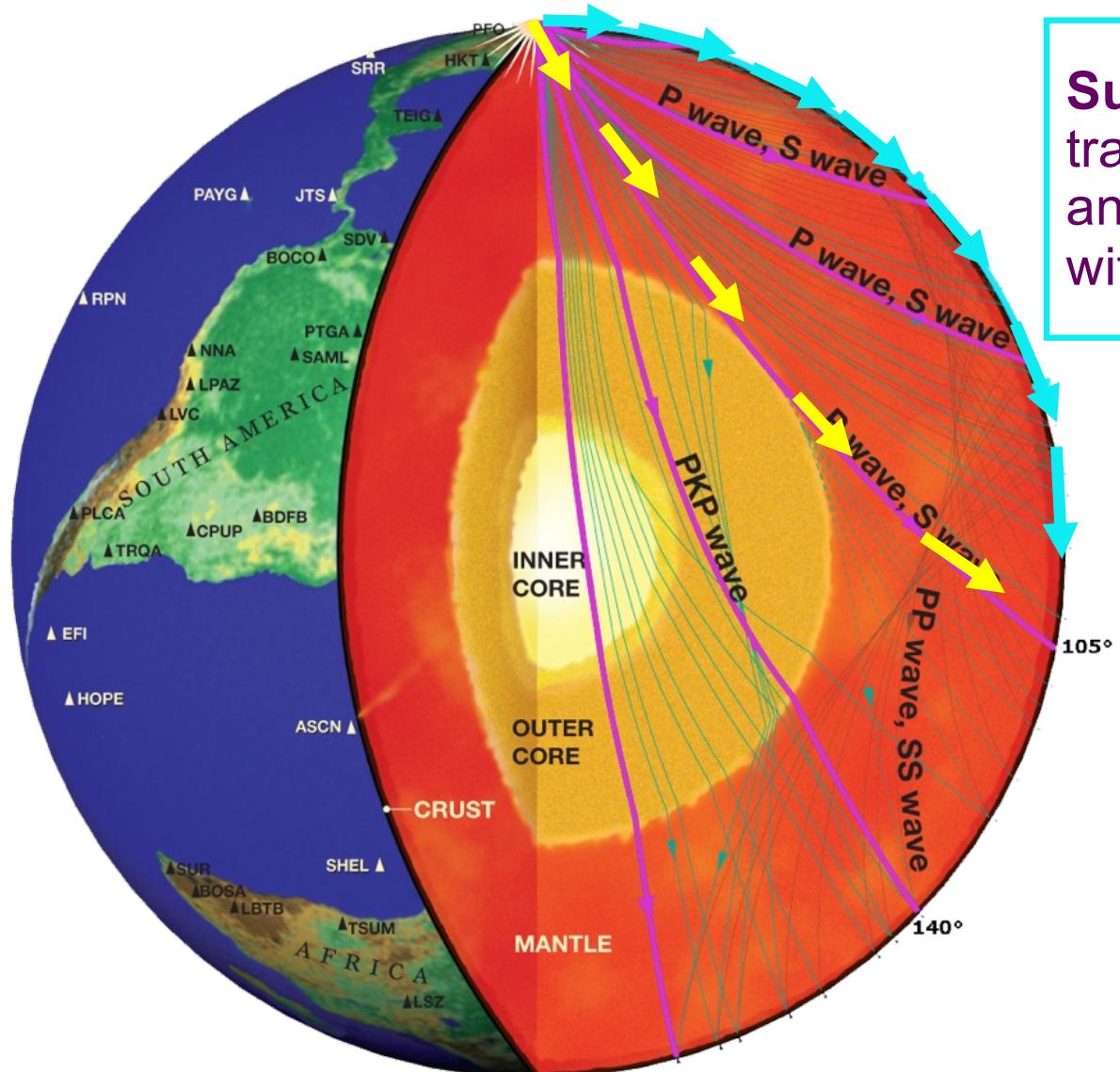
Body Waves



Surface Waves



Body Waves and Surface Waves



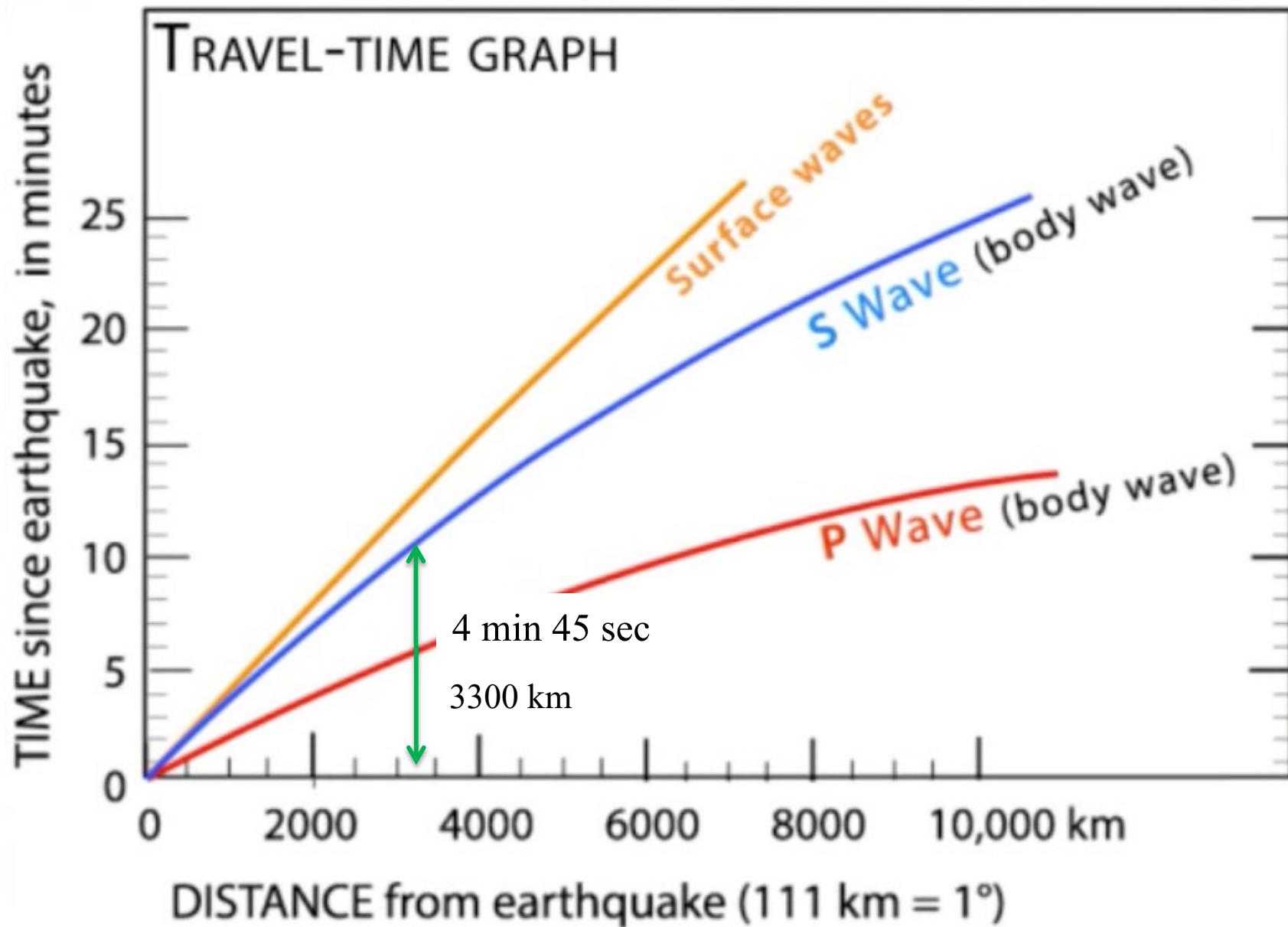
Surface waves

travel along Earth's surface, and decrease in amplitude with depth

Body waves

(P and S) radiate in all directions and travel inside Earth

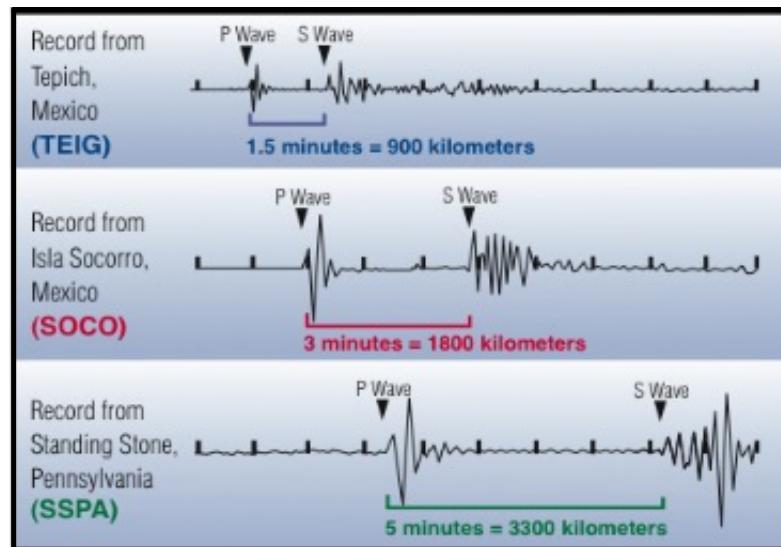
Distance of earthquake from seismometer



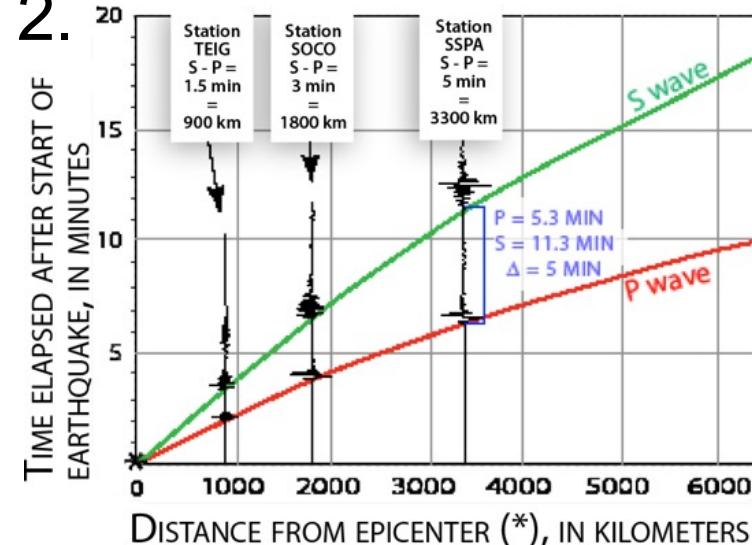
determined from S arrival time - P arrival time

Locating an earthquake...the basics

1.



2.



3.

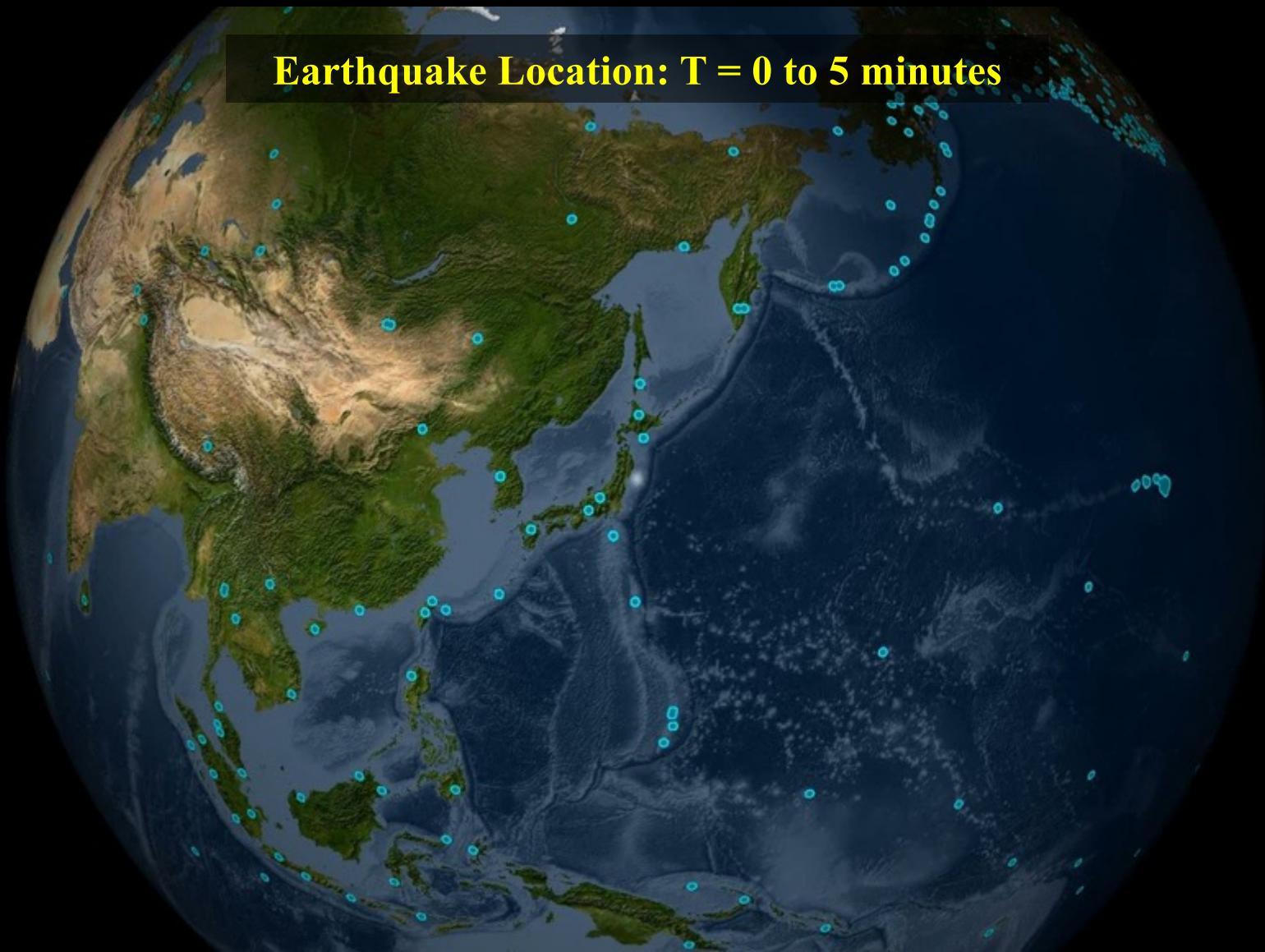
1. Determine distance of EQ from three seismic stations by calculating the S minus P arrival times.

2. Plot on the travel-time graph.

3. Intersection of the circles gives the location.



Earthquake Location: T = 0 to 5 minutes



Earthquake Rupture Complexity

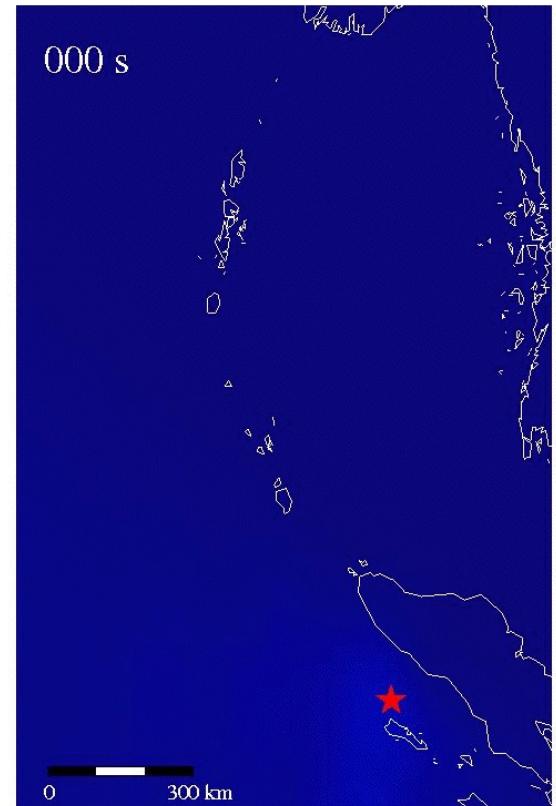
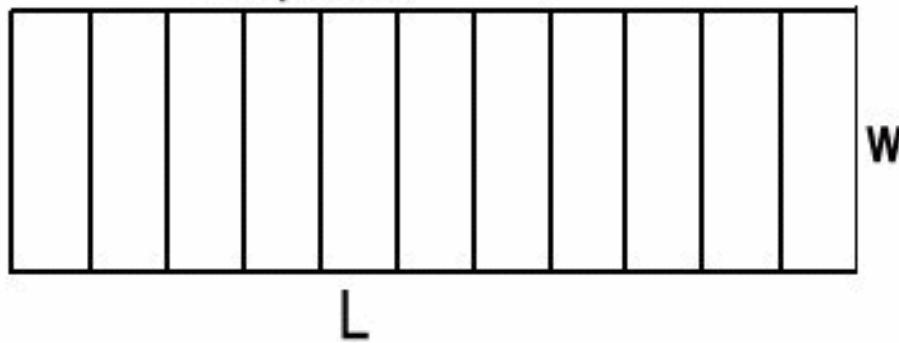
Great Earthquakes ($M \geq 8$)

- **Shake for a long time (10s sec to 2-3 minutes)**
- **Rupture for 100s miles**

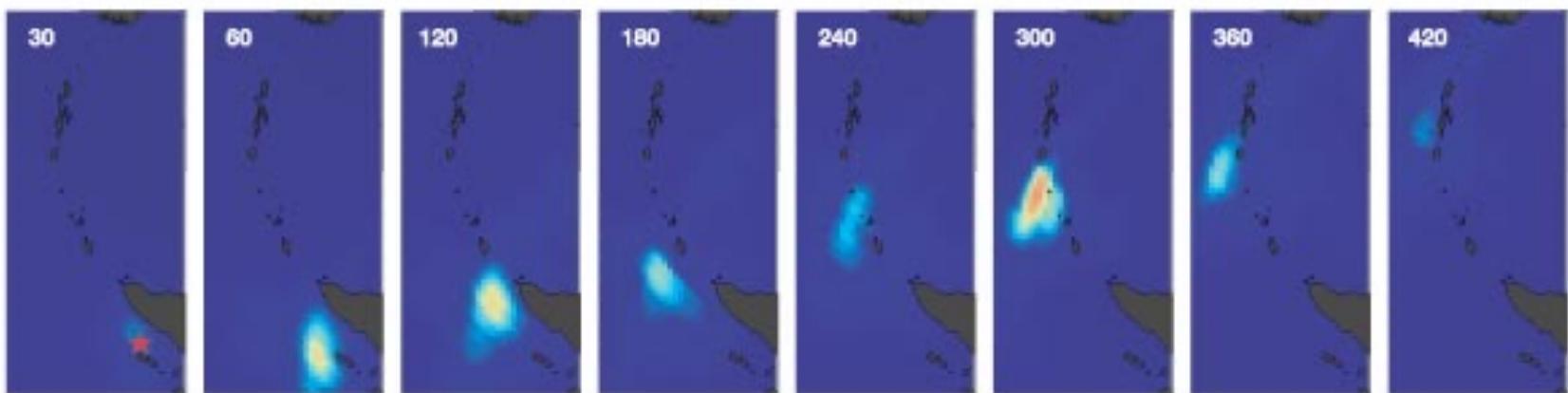
2004 Sumatra earthquake

Haskell Line Source Dislocation Source

Rupture ---->

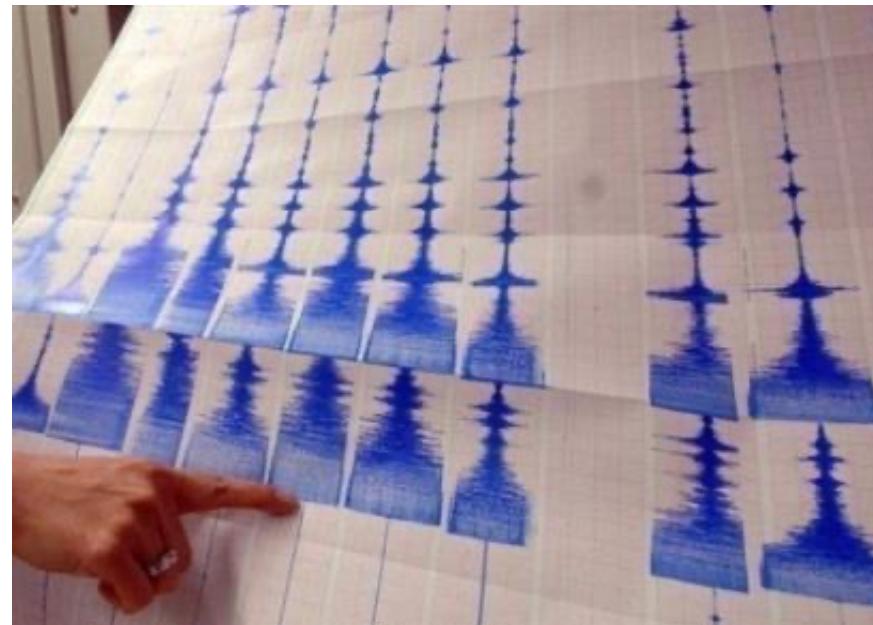


Energy Release imaged by Japan HINET Array



Earthquake Magnitude & Energy

M7, Papua New Guinea earthquake



TAKE 2

Myth-conceptions

Dispelling myths,
misconceptions, &
misunderstandings
about Earth science



INCORPORATED RESEARCH INSTITUTIONS FOR SEISMOLOGY

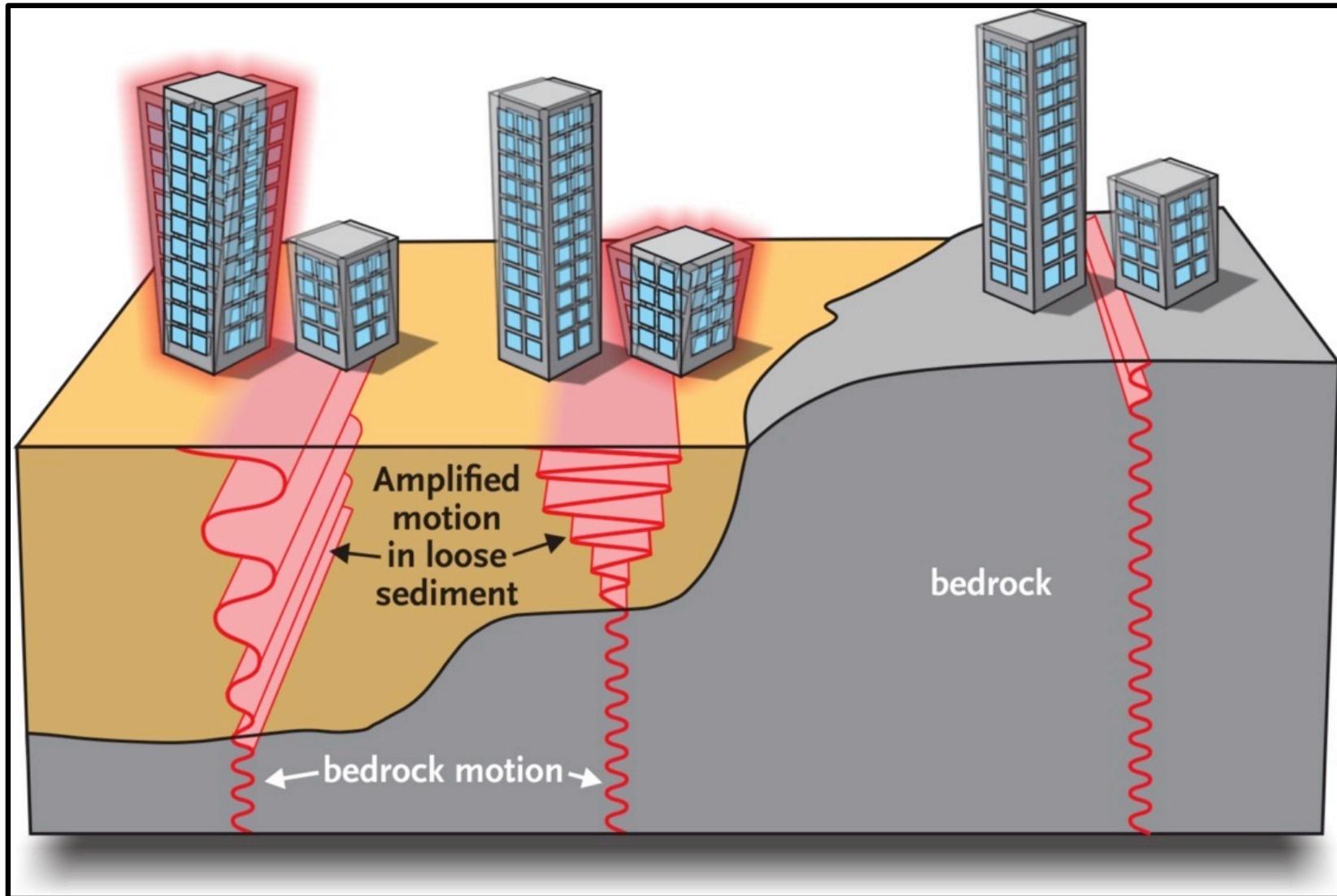


Earthquake Magnitude vs. Earthquake Intensity

Fixed value vs. variable

www.iris.edu/earthquake

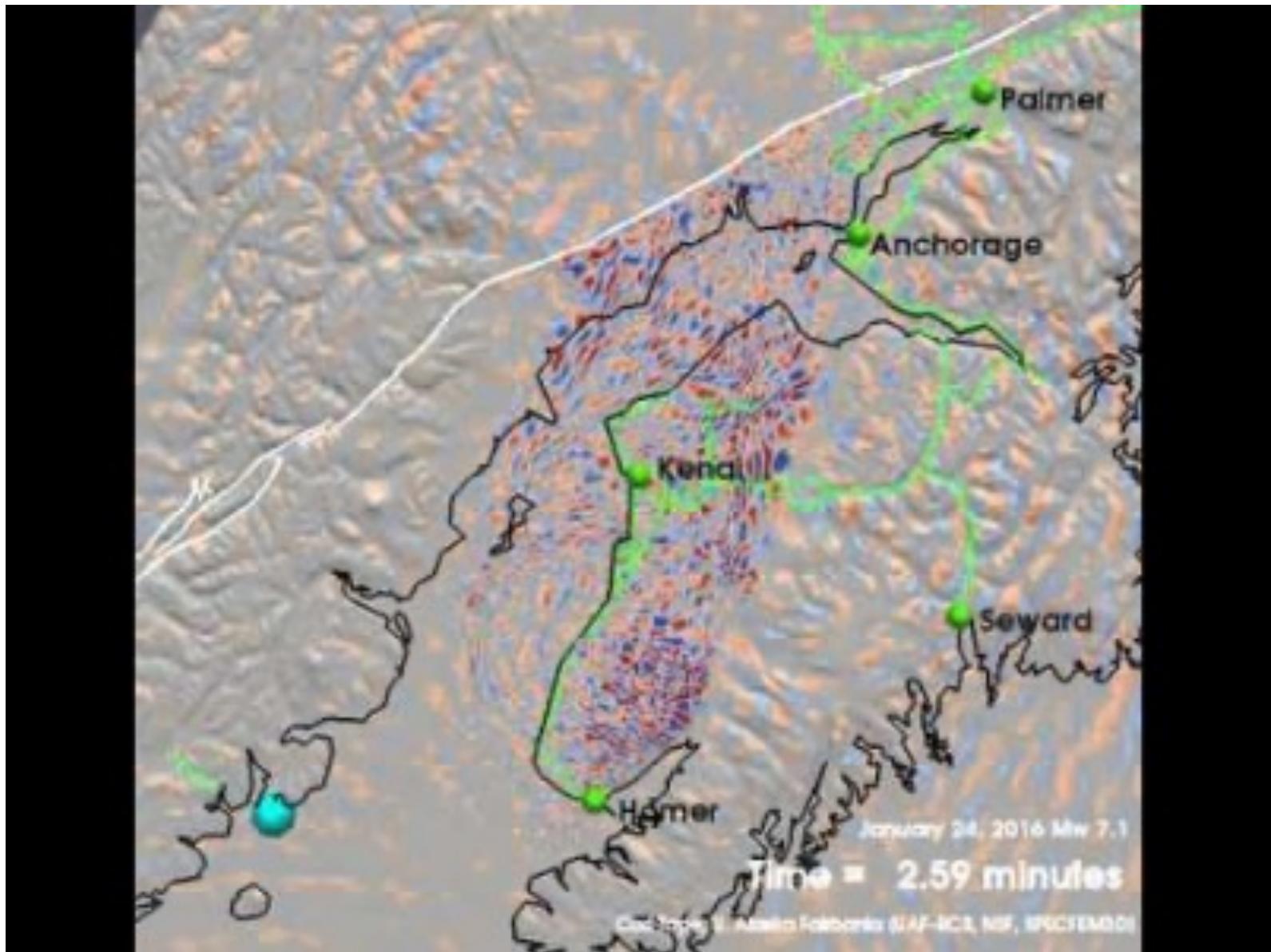
Ground-Shaking Amplification



Seismic waves are amplified as they pass from bedrock into basins filled with sedimentary rock.

Site Effects Related to soils and topography

M7.1 Pedro Bay Earthquake (animation by Carl Tape)



Measuring Earthquake size

- **Historical – Used macroseismic information**
 - Fatalities
 - Maximum shaking
 - Area of intense shaking
- **Did not correlate well from one quake to the next because damage depended on**
 - "True" size (i.e., magnitude)
 - Distance from the epicenter
 - Building design
 - Surface material (rock or dirt) beneath buildings
 - Proximity to populated regions

Modified Mercalli Intensity Scale

CIIM Intensity	People's Reaction	Furnishings	Built Environment	Natural Environment
I	Not felt			Changes in level and clarity of well water are occasionally associated with great earthquakes at distances beyond which the earthquakes felt by people.
II	Felt by a few.	Delicately suspended objects may swing.		
III	Felt by several; vibration like passing of truck.	Hanging objects may swing appreciably.		
IV	Felt by many; sensation like heavy body striking building.	Dishes rattle.	Walls creak; window rattle.	
V	Felt by nearly all; frightens a few.	Pictures swing out of place; small objects move; a few objects fall from shelves within the community.	A few instances of cracked plaster and cracked windows with the community.	Trees and bushes shaken noticeably.
VI	Frightens many; people move unsteadily.	Many objects fall from shelves.	A few instances of fallen plaster, broken windows, and damaged chimneys within the community.	Some fall of tree limbs and tops, isolated rockfalls and landslides, and isolated liquefaction.
VII	Frightens most; some lose balance.	Heavy furniture overturned.	Damage negligible in buildings of good design and construction, but considerable in some poorly built or badly designed structures; weak chimneys broken at roof line, fall of unbraced parapets.	Tree damage, rockfalls, landslides, and liquefaction are more severe and widespread with increasing intensity.
VIII	Many find it difficult to stand.	Very heavy furniture moves conspicuously.	Damage slight in buildings designed to be earthquake resistant, but severe in some poorly built structures. Widespread fall of chimneys and monuments.	
IX	Some forcibly thrown to the ground.		Damage considerable in some buildings designed to be earthquake resistant; buildings shift off foundations if not bolted to them.	
X			Most ordinary masonry structures collapse; damage moderate to severe in many buildings designed to be earthquake resistant.	

USGS ShakeMap

2010 M_w 8.8 Maule, Chile

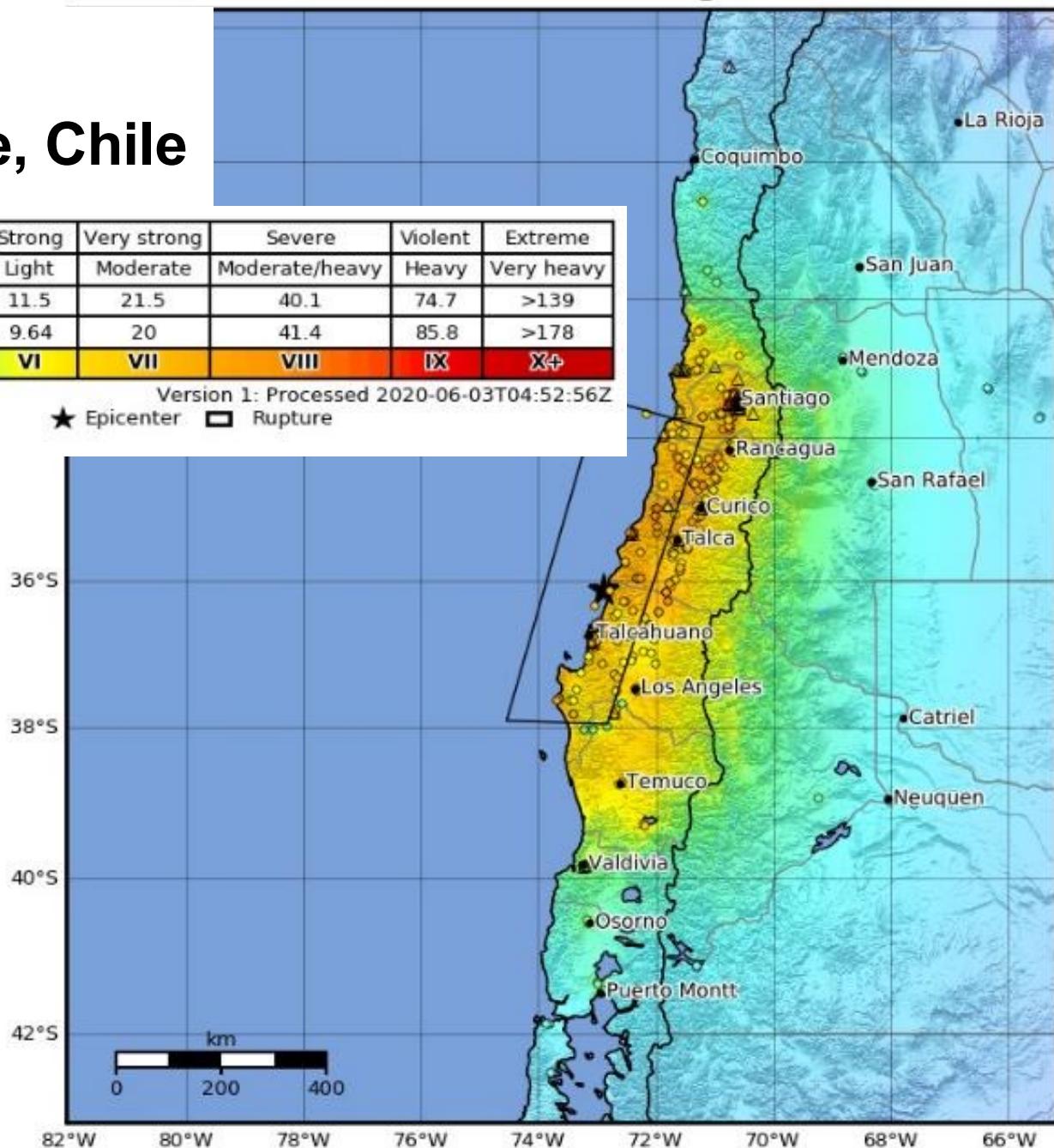
SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	None	None	None	Very light	Light	Moderate	Moderate/heavy	Heavy	Very heavy
PGA(%g)	<0.0464	0.297	2.76	6.2	11.5	21.5	40.1	74.7	>139
PGV(cm/s)	<0.0215	0.135	1.41	4.65	9.64	20	41.4	85.8	>178
INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+

Scale based on Worden et al. (2012)

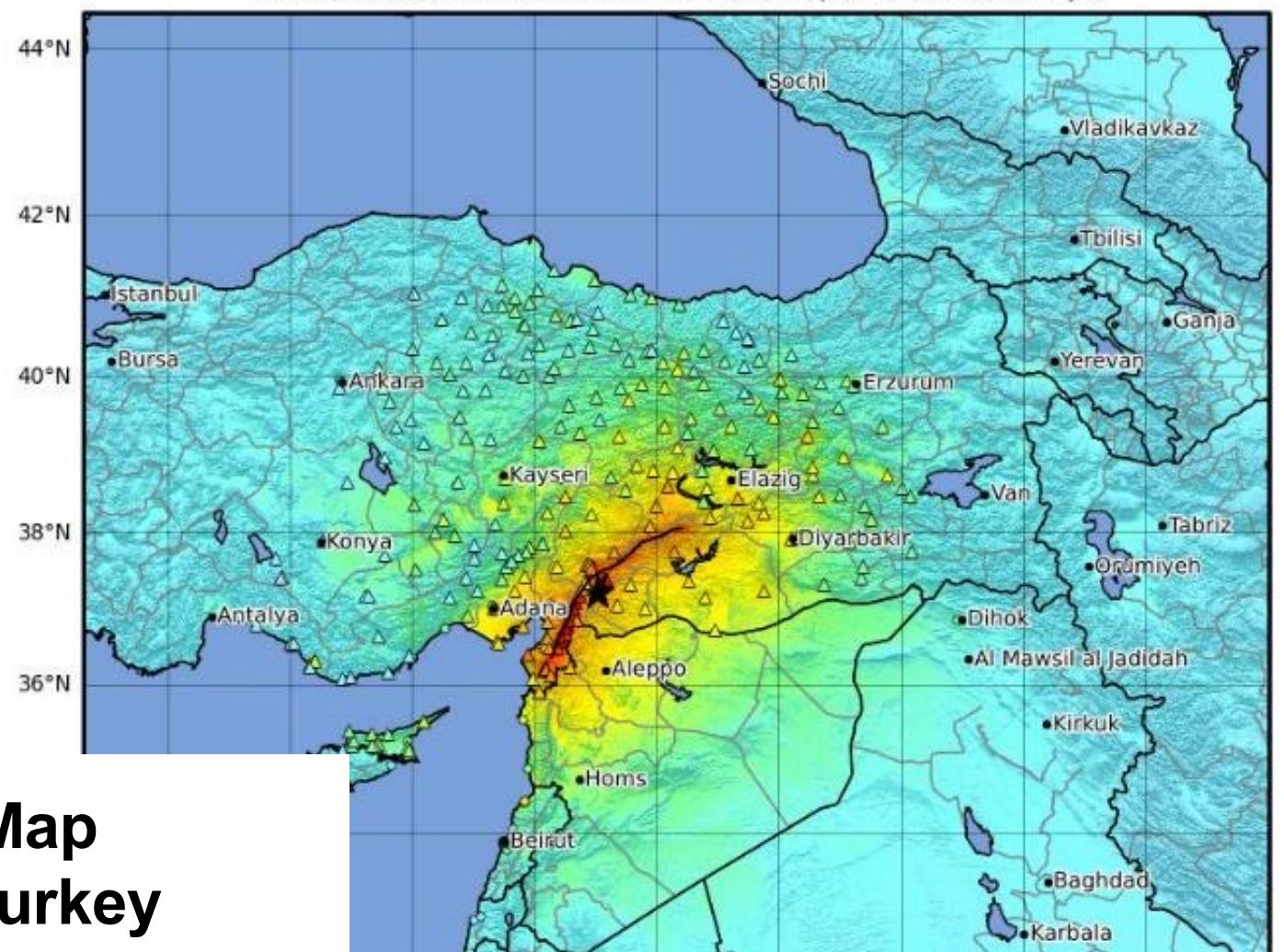
△ Seismic Instrument ○ Reported Intensity

Version 1: Processed 2020-06-03T04:52:56Z

★ Epicenter □ Rupture



Macroseismic Intensity Map USGS
 ShakeMap: 25 km ENE of Nurdağı, Gaziantep, TR
 Feb 06, 2023 01:17:34 UTC M7.8 N37.23 E37.01 Depth: 10.0km ID:us6000jllz



USGS ShakeMap 2023 M_w 7.8 Turkey

SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	None	None	None	Very light	Light	Moderate	Moderate/heavy	Heavy	Very heavy
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Version 1: Processed 2020-06-03T04:52:56Z

★ Epicenter □ Rupture

E 42°E 44°E 46°E 48°E

Earthquake Magnitude

General form:

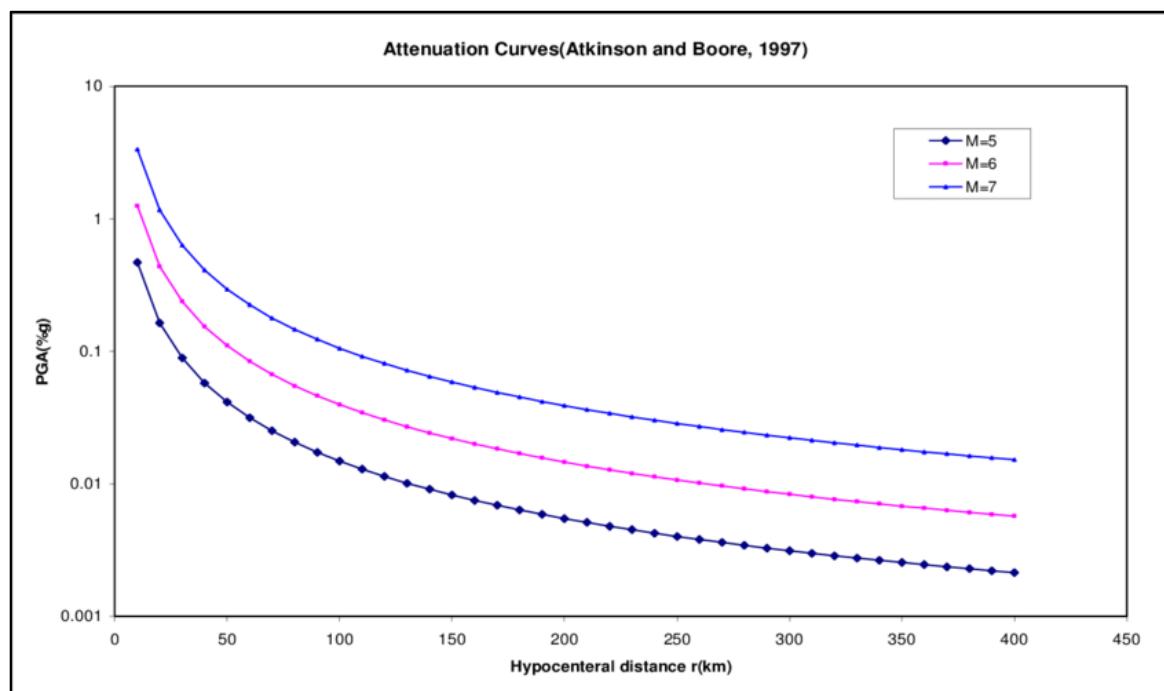
$$M = \log(A/T) + F(h, \Delta) + C$$

Amplitude

Period

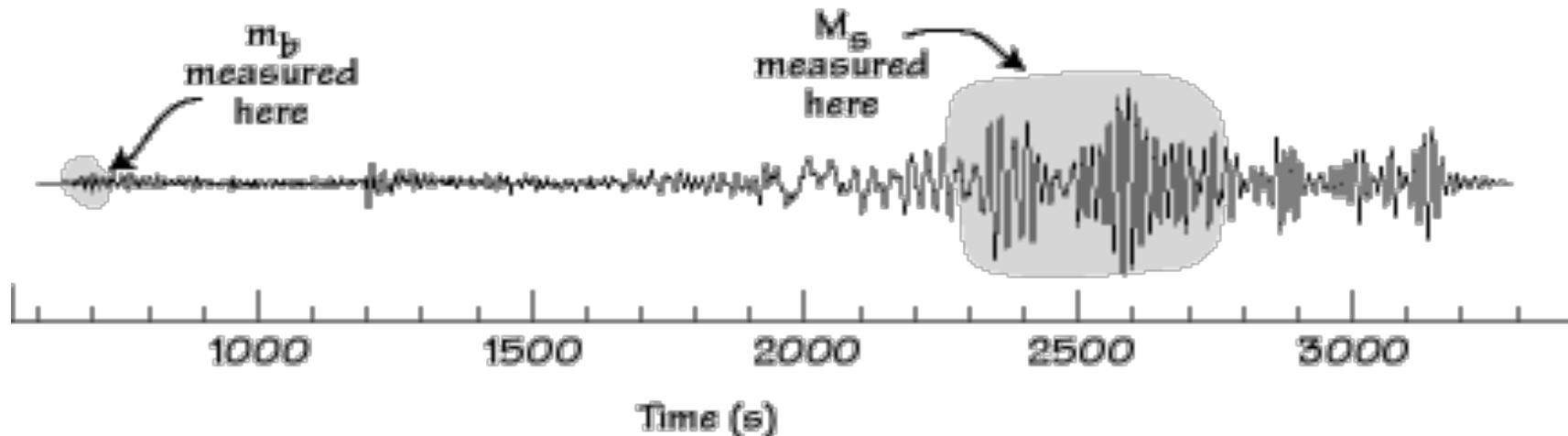
Regional scaling factor

Correction for depth
and distance



Richter and Gutenberg's Teleseismic (distant) magnitudes

Body wave Magnitude mb and Surface wave and Ms



$$mb = \log (A/T) + Q(D, h)$$

T: period (secs), $0.1 \leq T \leq 3.0$

A: P wave amplitude (microns)
(not necessarily the maximum)

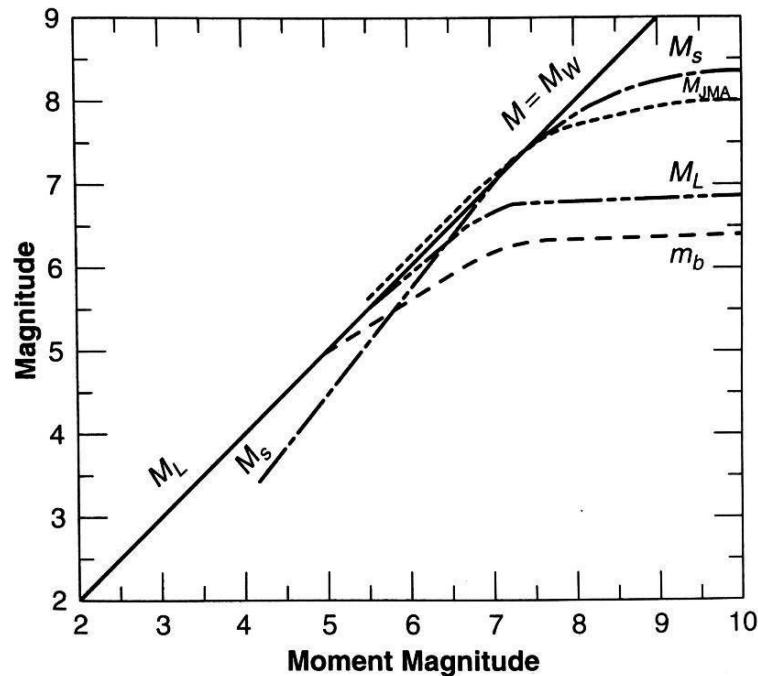
Q: scale factor ($D \geq 5^\circ$)

$$Ms = \log (A/T) + 1.66 \log D + 3.3$$

A: maximum amplitude (microns)
vertical component of the
surface wave within the period
range $18 \leq T \leq 22$.
D: $20^\circ \leq D \leq 160^\circ$.

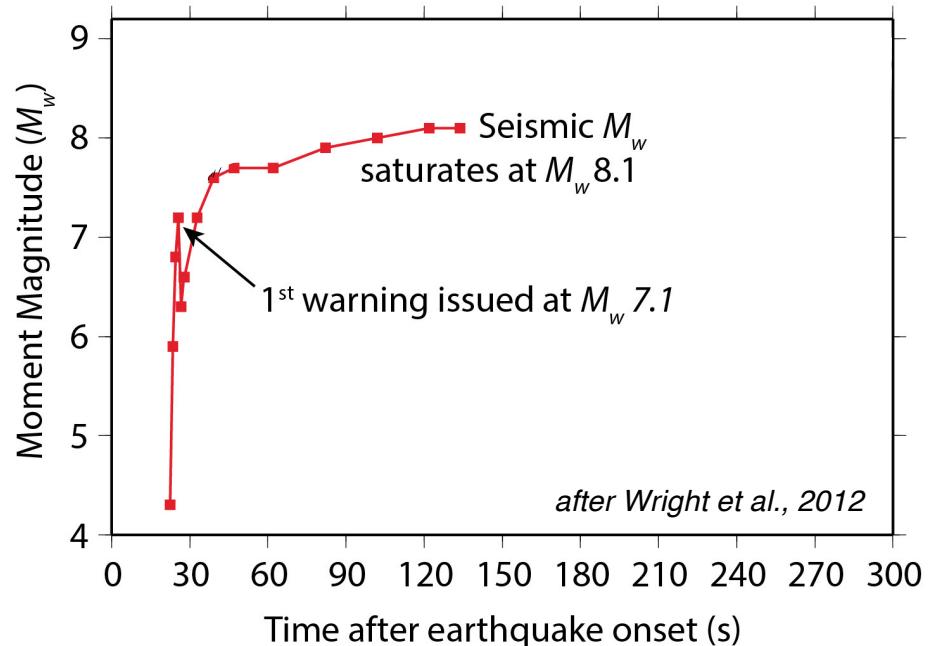
No depth corrections!

Saturation!



Idriss, 1985

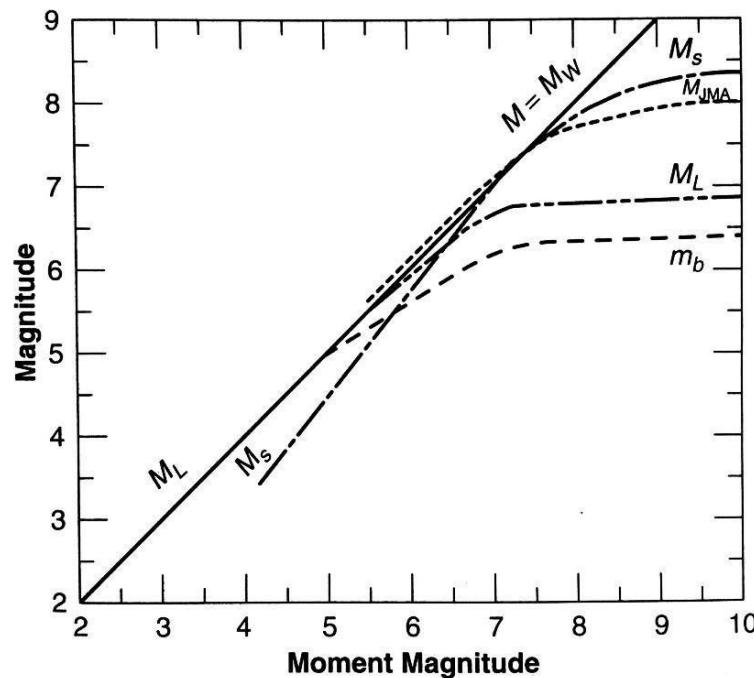
2011 Tohoku-Oki (Japan) Earthquake



When does this happen?

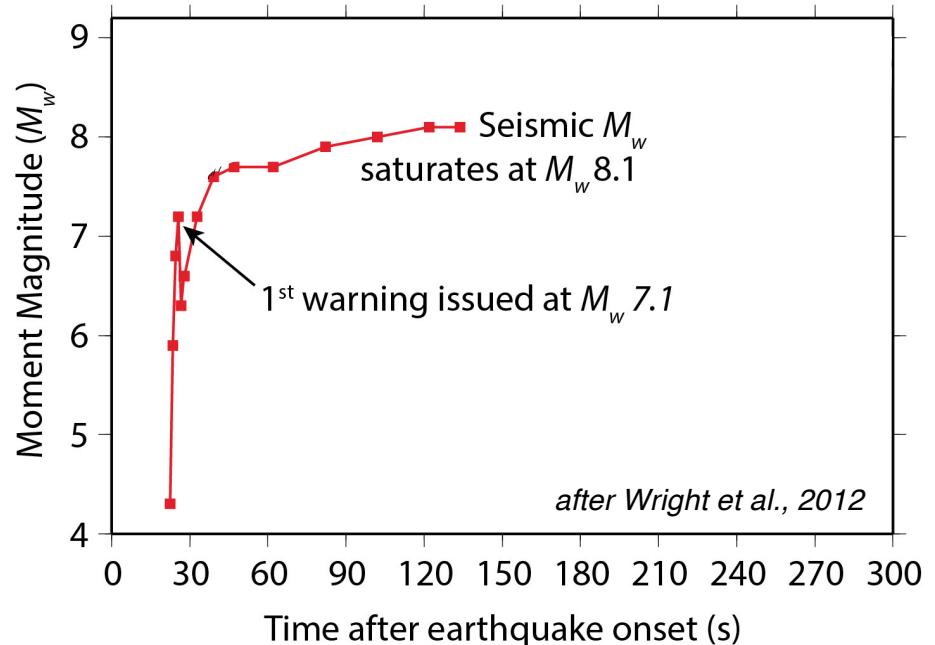
- 1) **Time window saturation** – when time window is less than duration of rupture (particularly affects m_b)
- 2) **Spectral saturation** – when wavelength too short to see entire rupture (affects m_b , M_L , and M_s)

Saturation!



Idriss, 1985

2011 Tohoku-Oki (Japan) Earthquake



How do we overcome this problem?

- Examine longer period waves
- M_w , M_{wp} , Mantle magnitude (M_m), Centroid Moment Tensor (CMT) – moment magnitude
- GNSS/GPS data (next talk)

Types of Magnitude Scales

		Period Range
M_L	Local magnitude (California)	regional S & surface waves 0.1-1 sec
M_j	JMA (Japan Meteorol. Agency)	regional S & surface waves 5-10 sec
m_b	Body wave magnitude	teleseismic P waves 1-5 sec
M_s	Surface wave magnitude	teleseismic surface waves 20 sec

The methods below overcome the effects of saturation:

M_{wp}	P-wave moment magnitude	teleseismic P waves	10-60 sec
M_w	Moment magnitude	teleseismic surface waves	> 200 sec
M_m	Mantle magnitude	teleseismic surface waves	> 200 sec

Moment magnitude (Mw)

$$Mw = (2/3) \log_{10} (M_0) - 16.1$$

- Introduced in 1979 by Hanks and Kanamori
- Based on source parameter M_0 and is not frequency dependent, **does NOT** saturate
- Based on earthquake **energy release**
- Related to fault slip and not ground shaking
- Used to estimate the magnitude of large earthquakes
- Very useful for **tsunami modeling**

Moment Magnitude

Understanding moment magnitude
...or what bumped the Richter scale?



IRIS
Institutes for the
Planets and the
Sun

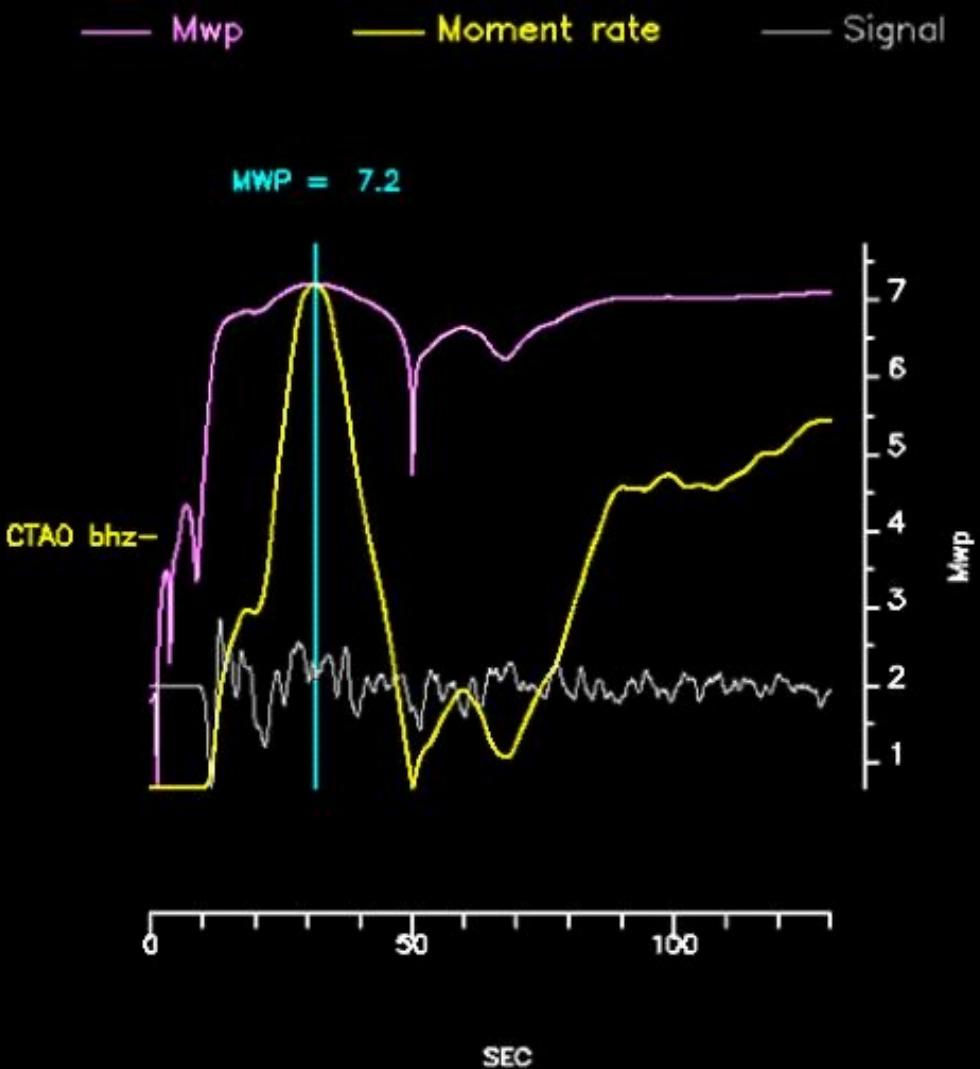


Mwp Magnitude

- Moment magnitude based on **initial long-period P-waves**
- Developed by S. Tsuboi and others (1995)
- **Empirical estimate of the moment magnitude**
 - integrate the vertical velocity from a seismogram
- **Accurate results within 3-4 minutes of OT (P, pP arrivals)**
- Primary initial **magnitude estimate at PTWC for M>6**
(replaced Ms)
- **Subject to site and path effects, source complexity, contamination from other large earthquakes**

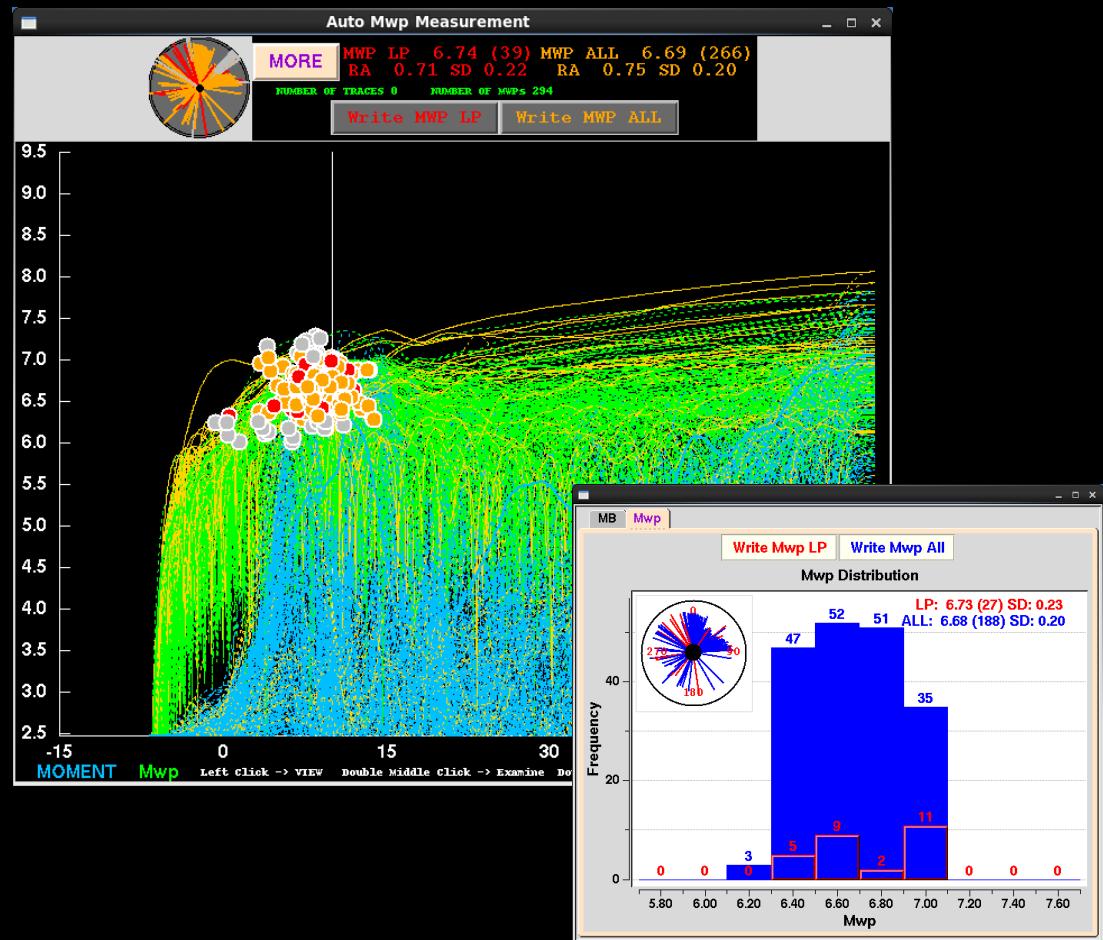
M_{wp} [Tsuboi et al., 1995;1999]

- Double integration of $v(t)$
 $v(t) \rightarrow M_o(t) \rightarrow M_w(t)$
- Peak $M_w(t) \rightarrow M_{wp}$
- Fast; less prone to saturation



M_{wp} [Tsuboi et al., 1995;1999]

- Double integration of $v(t)$
 $v(t) \rightarrow Mo(t) \rightarrow Mw(t)$
- Peak $Mw(t) \rightarrow Mwp$
- Fast; less prone to saturation



TWC magnitude determination – PTWC (5 primary)

- **ML** - local magnitude, HI/PR/USVI ($2 < M < 6$)
- **mb** - body wave magnitude, largest P-waves ($4 < M < 6.5$)
- **Mm** - mantle wave magnitude, 50-400s surface waves ($M > 6$)
- **Mwp** – estimate of moment magnitude from integrated P-waves ($5 < M < 8$)
- **W-phase** - long-period phase, gives stable results, does not saturate, but takes $\sim 20\text{-}25$ minutes ($M > 5$)

TWC magnitude determination – PTWC (5 primary)

- Measurements are not precise.
Different magnitudes result from using different scales/methods and datasets.
- **Many methods saturate for large earthquakes**
- **Mo → EQ energy → "true" EQ size but takes time**
- **TWC's rely on Mwp for preliminary message products, method can still underestimate great earthquake sizes.**
- **WCMT provides fault geometry and authoritative estimate of Mw but takes time (~15-25 minutes)**

Thank You

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