



UNESCO-IOC / NOAA ITIC Training Program in Hawaii (ITP-TEWS Hawaii)  
TSUNAMI EARLY WARNING SYSTEMS  
AND THE PACIFIC TSUNAMI WARNING CENTER (PTWC) ENHANCED PRODUCTS  
TSUNAMI EVACUATION PLANNING AND UNESCO IOC TSUNAMI READY PROGRAMME  
15-26 September 2025, Honolulu, Hawaii

# **Tsunami Warning Center Operations: Real-Time Earthquake Detection and Fast Source Characterization – Locating Earthquakes & Estimating Magnitude**

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NOAA Pacific Tsunami Warning Center

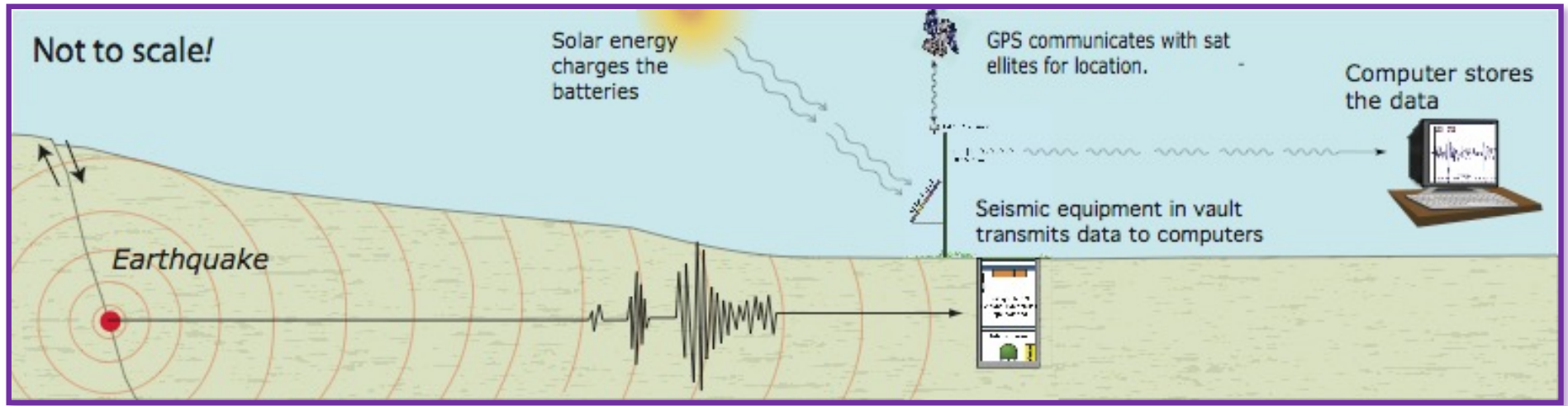


Pacific  
Community  
Communauté  
du Pacifique

# Locating Earthquakes

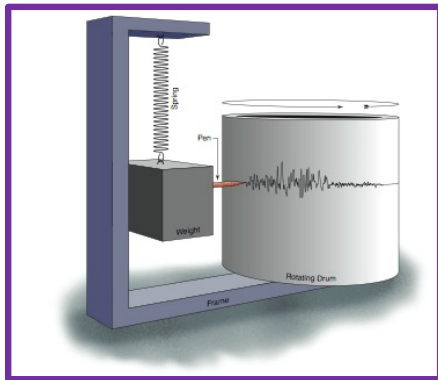
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# How do we measure earthquakes?

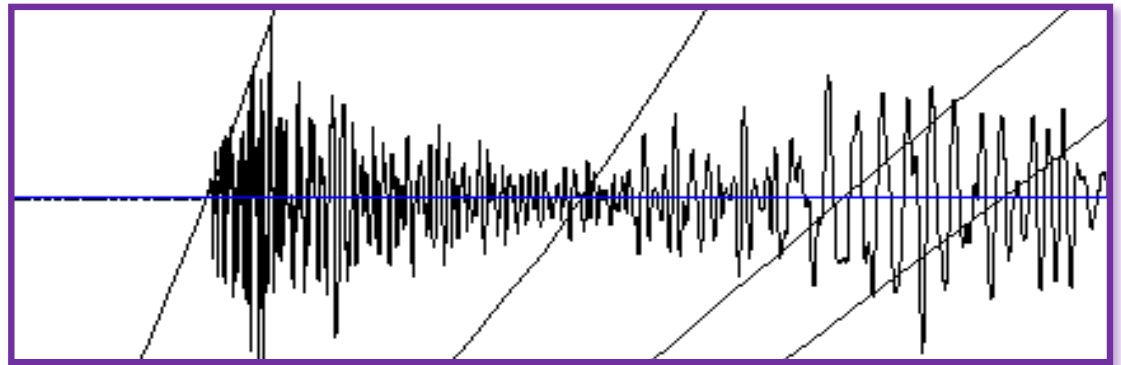


When an earthquake occurs, the seismic waves travel through the Earth to the seismic station where the information is transmitted to distant computers.

A **seismograph** detects and records earthquakes



A **seismogram** is the earthquake record



# Global Seismic Network

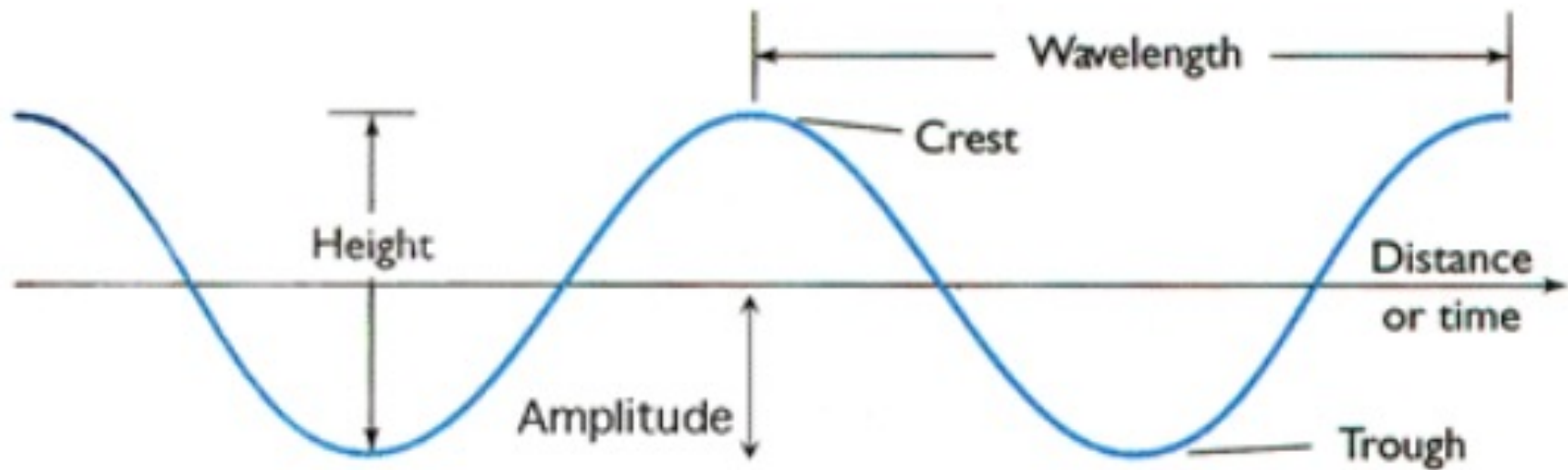


● Seismometers

PACIFIC TSUNAMI  
WARNING CENTER

PTWC monitors >800 seismic stations worldwide

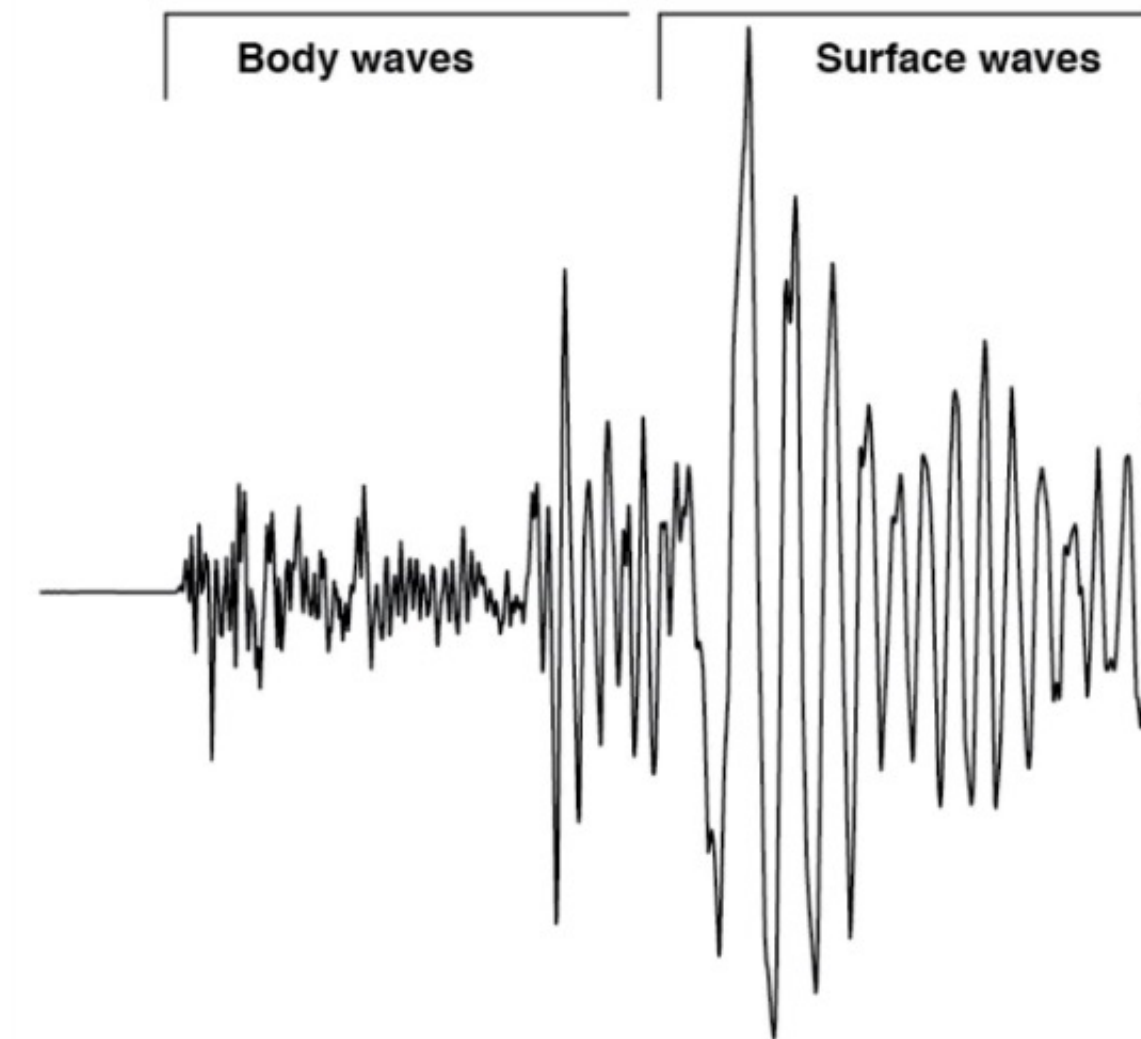
# Basic (Seismic) Wave Properties



- Wavelength - distance between successive crests of the wave
- Wave period - time it takes for the wave to travel one wavelength



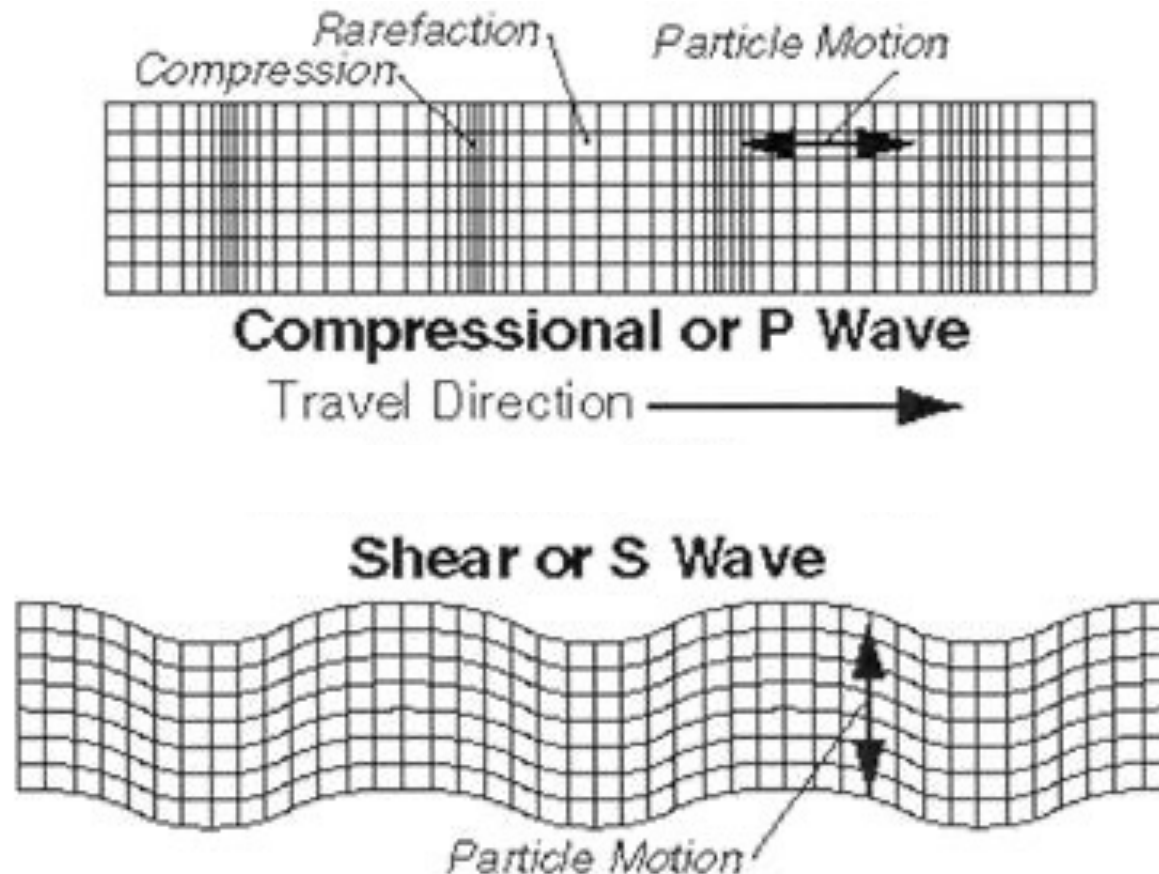
# Basic Types of Seismic Waves



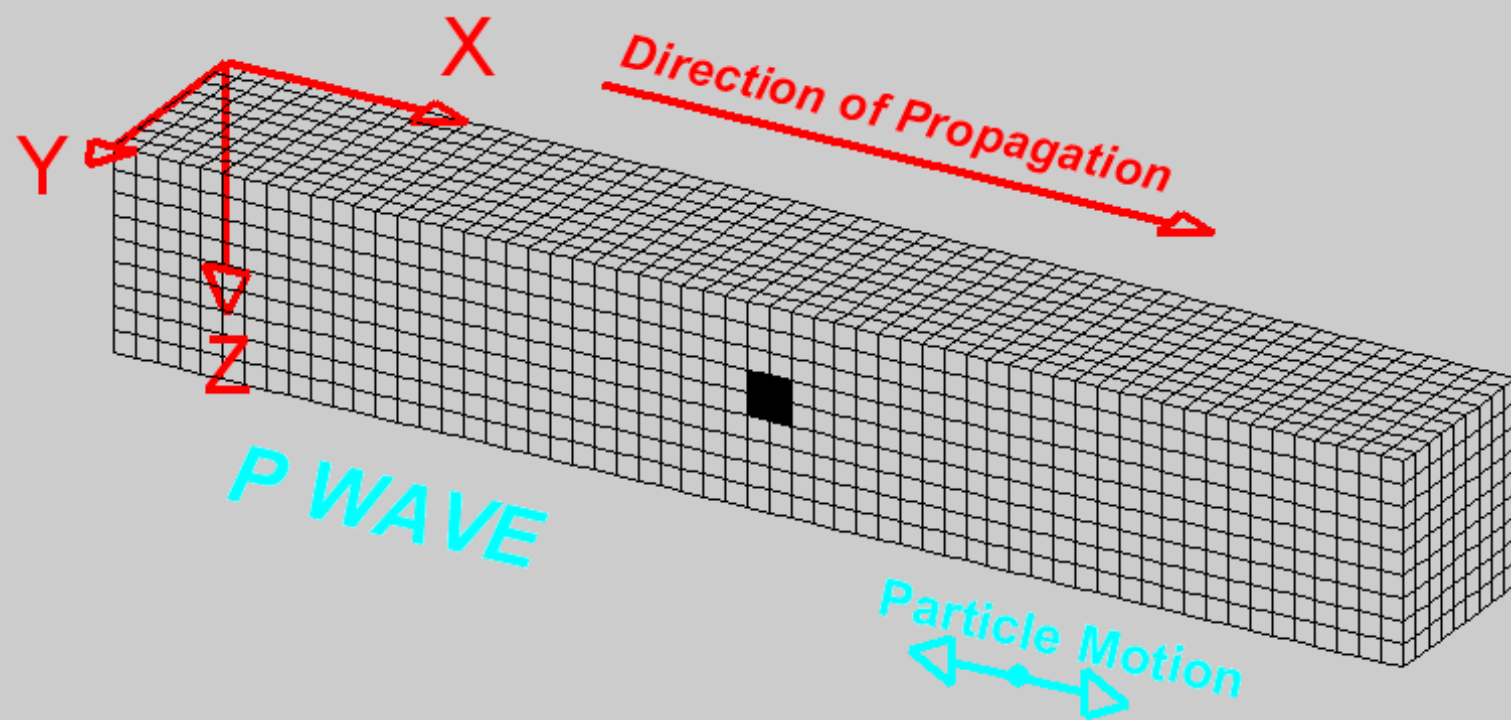
Surface waves arrive after the body waves. They have lower frequency and larger amplitude than body waves.

# Basic Types of Seismic Waves

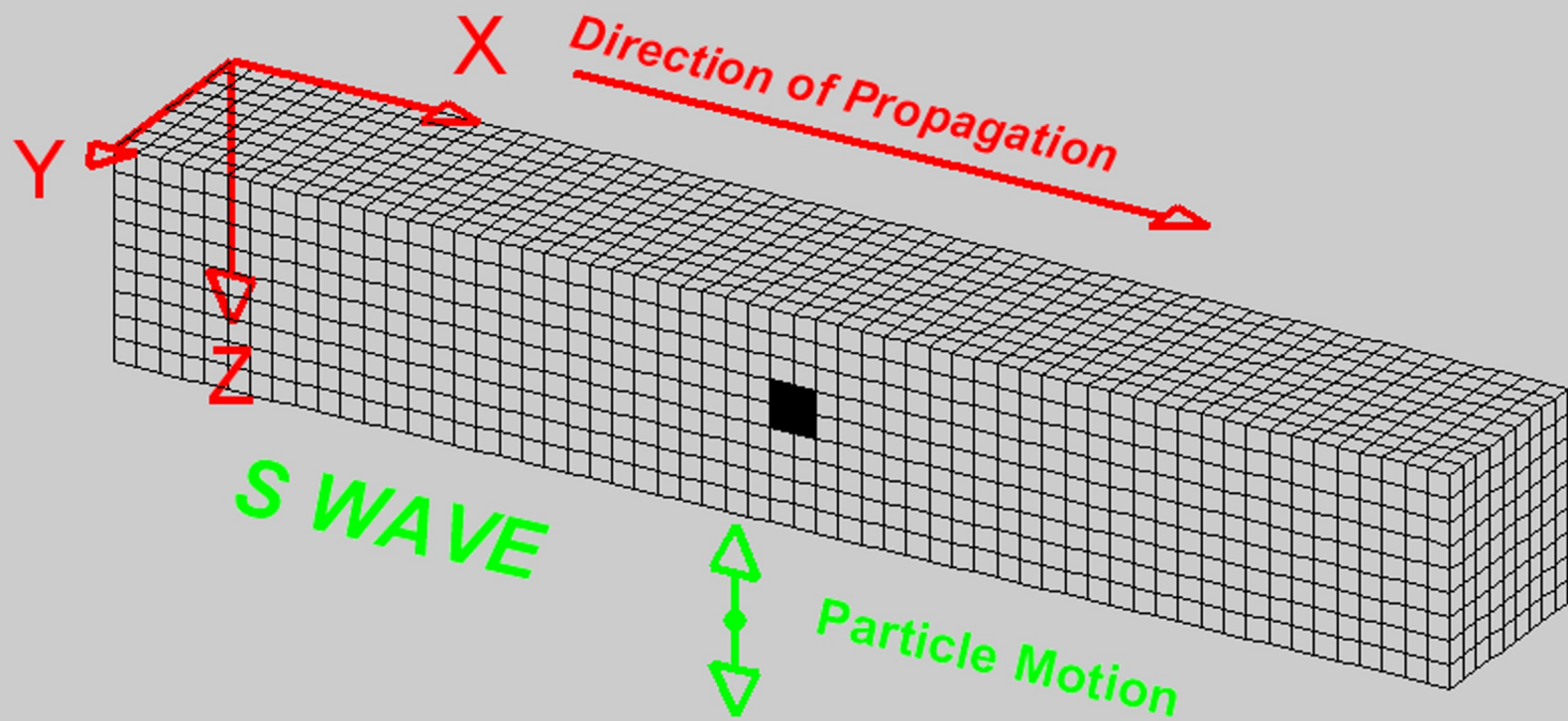
## P and S Waves (also called body waves)



P waves are faster and travel at speeds of **6-14 km/s**

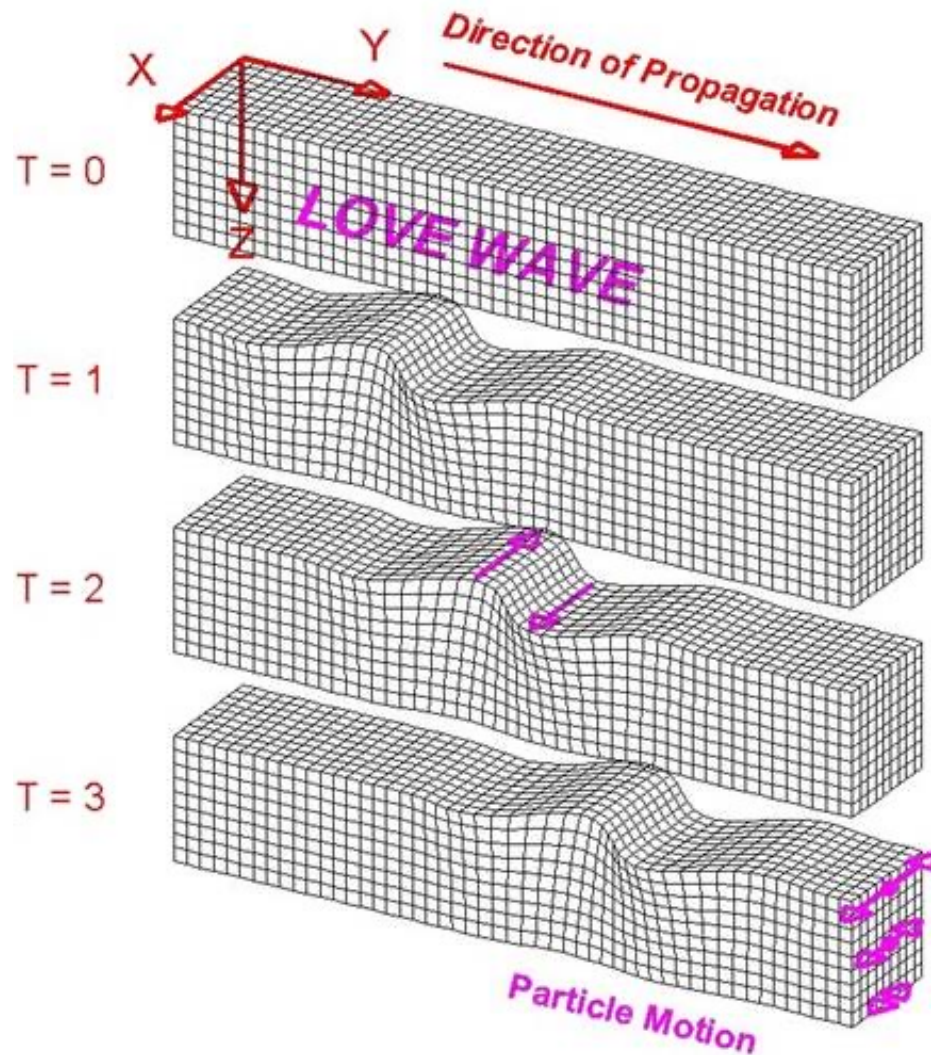




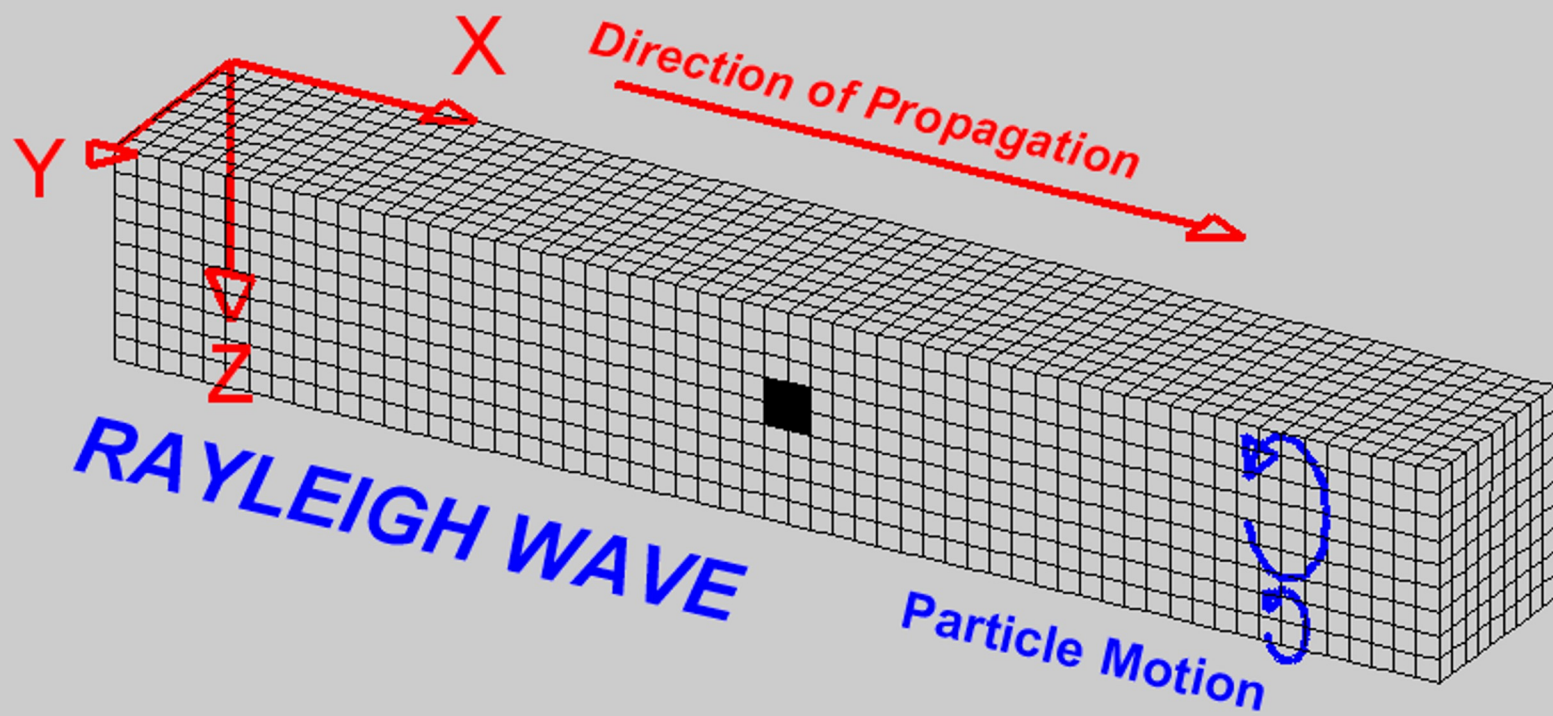


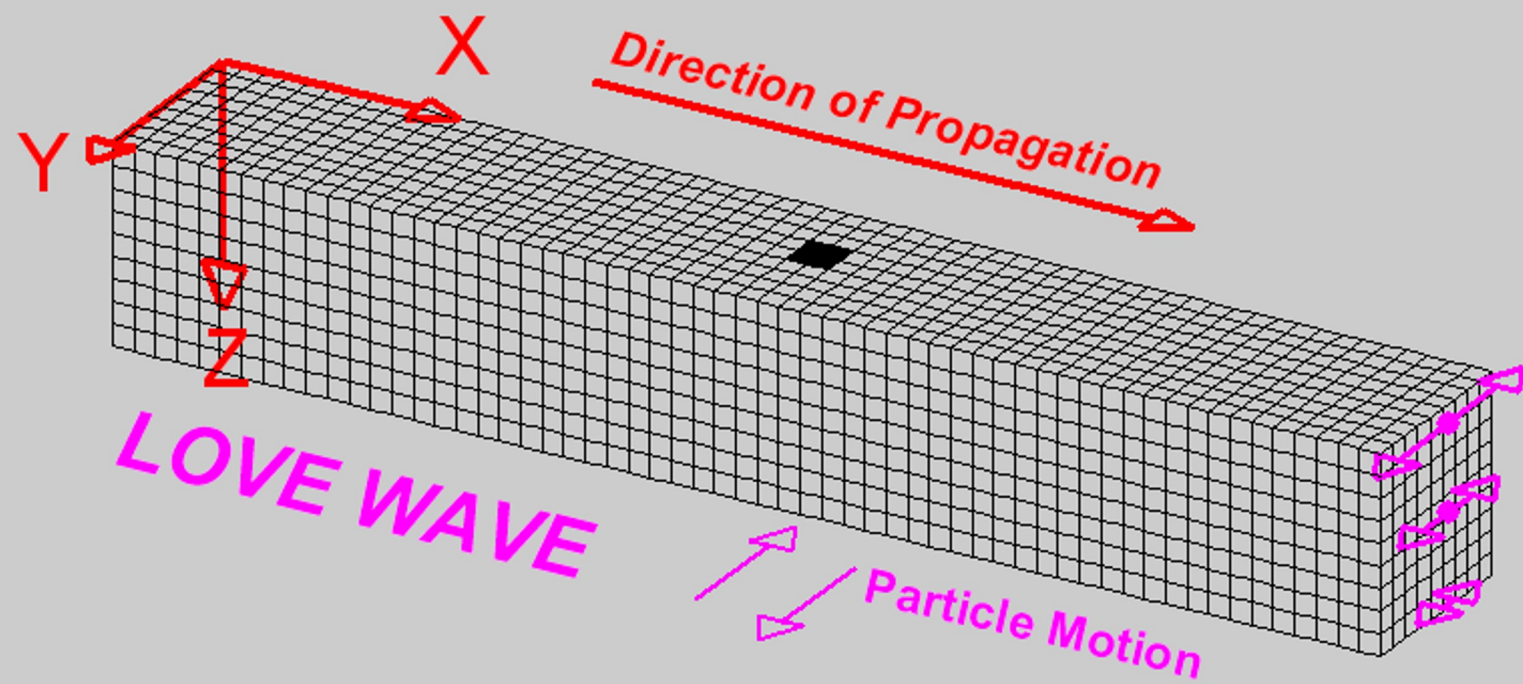
# Basic Types of Seismic Waves

## Surface Waves (do not travel through the earth)



Surface waves travel along the Earth's surface at speeds of  $\sim 3/\text{km/sec}$ , much slower than the body waves.

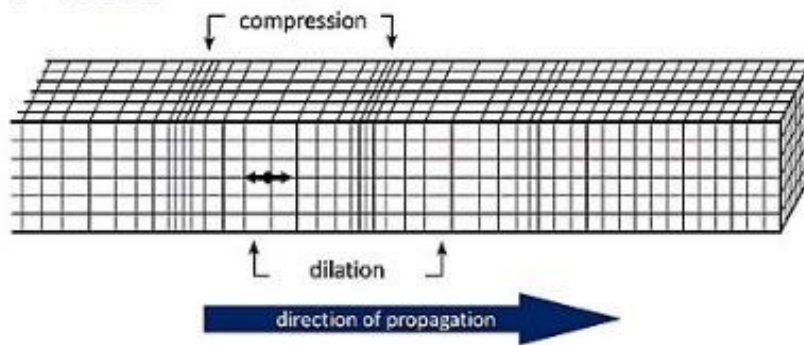




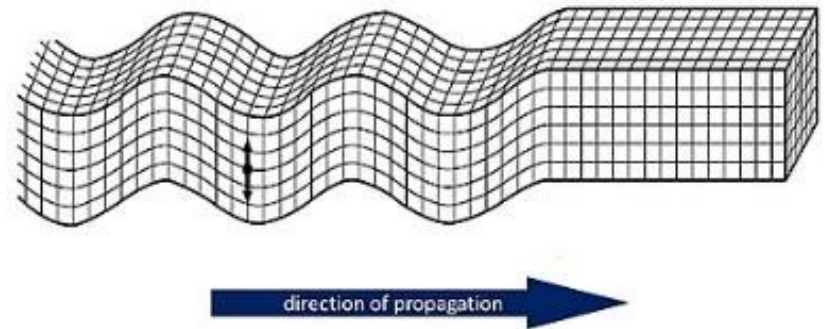


Body Waves

P Wave

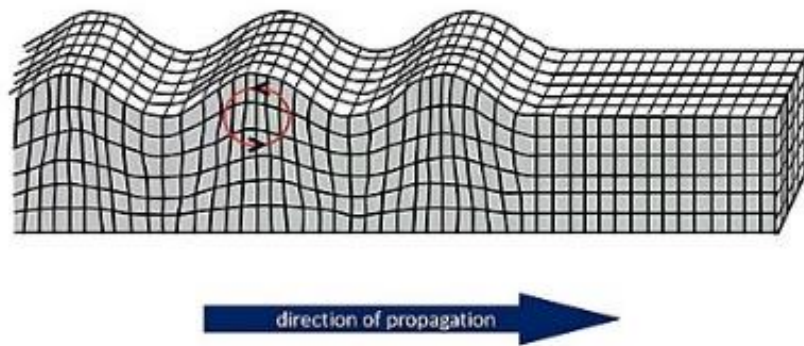


S Wave

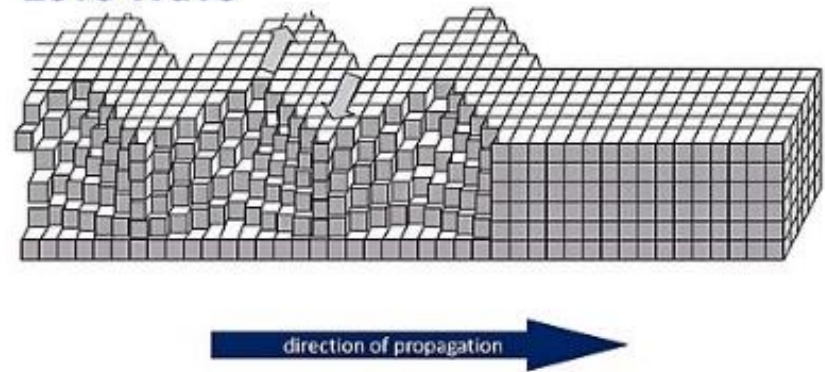


Surface Waves

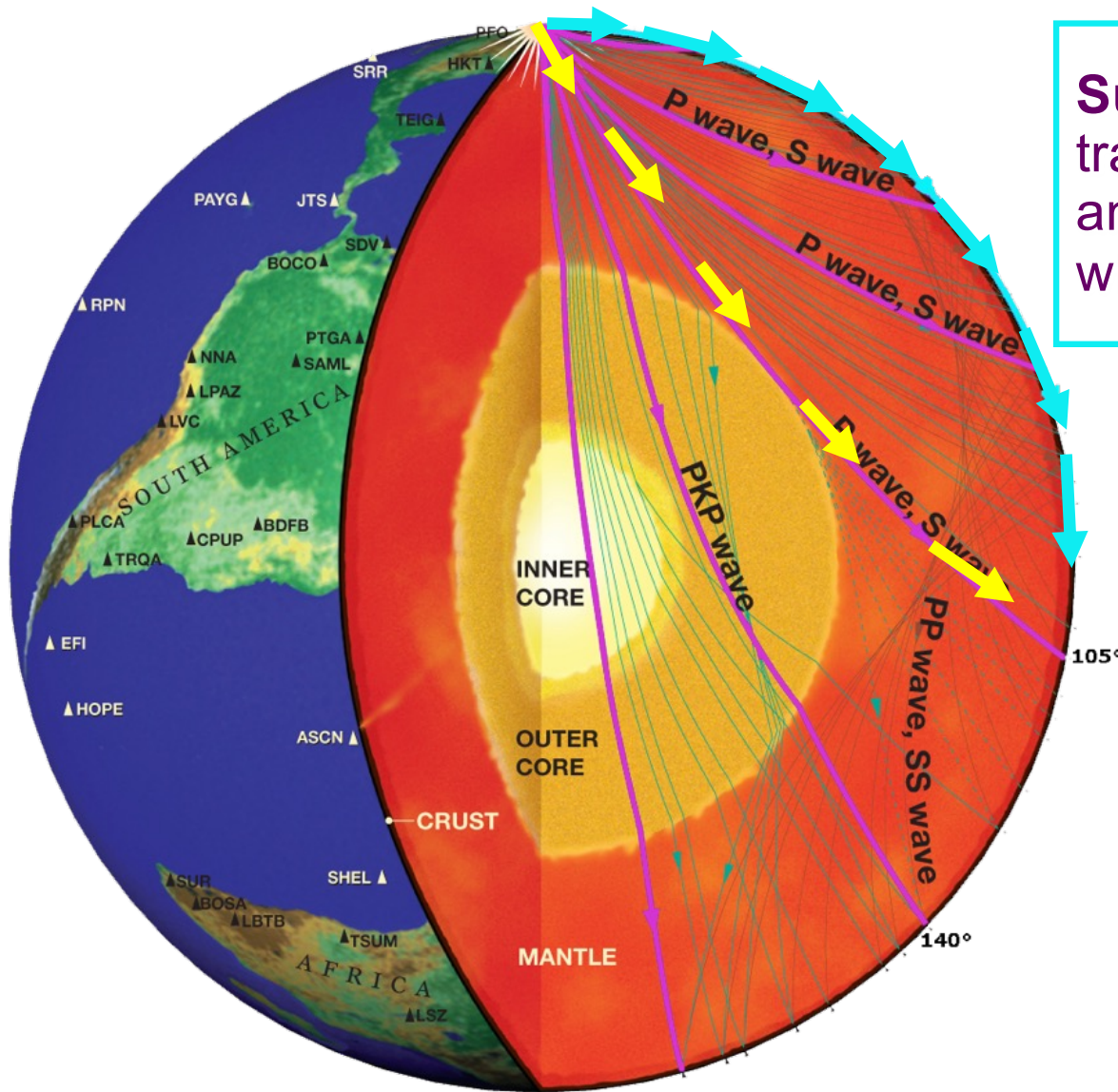
Rayleigh Wave



Love Wave



# Body Waves and Surface Waves

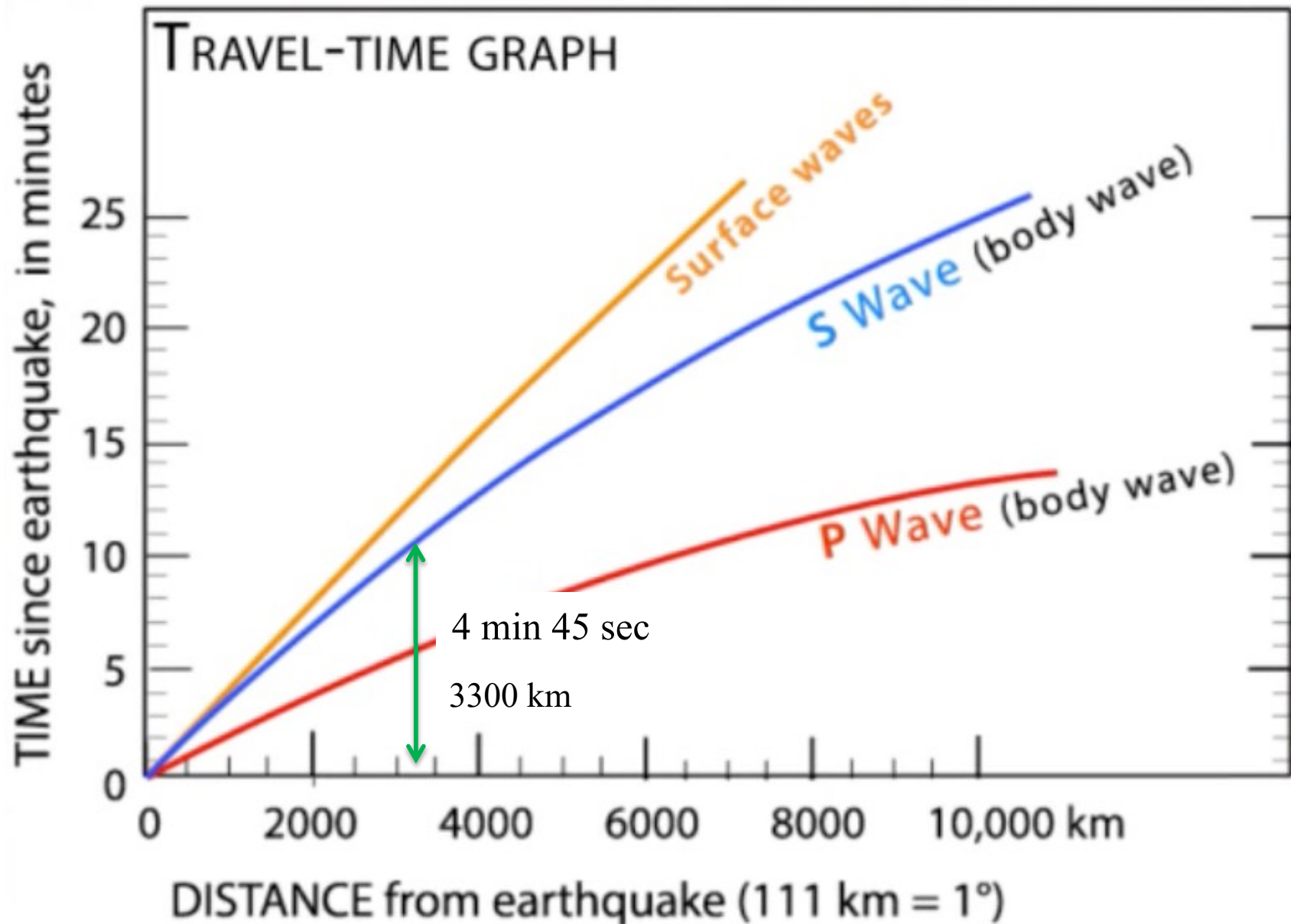


**Surface waves**  
travel along Earth's surface,  
and decrease in amplitude  
with depth

**Body waves**  
(P and S) radiate in  
all directions and  
travel inside Earth



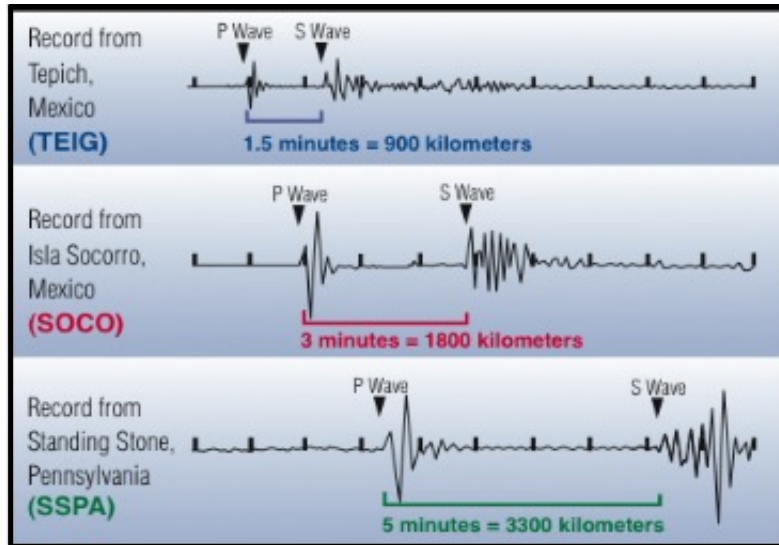
# Distance of earthquake from seismometer



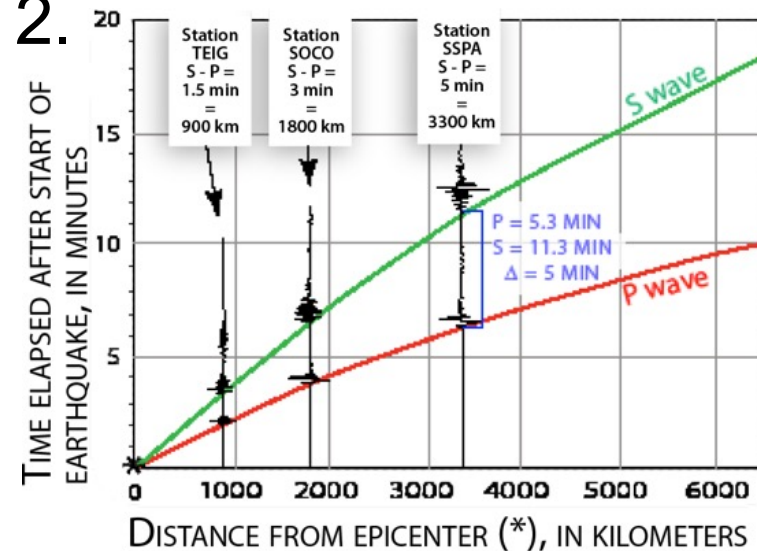
determined from S arrival time - P arrival time

# Locating an earthquake...the basics

1.



2.



1. Determine distance of EQ from three seismic stations by calculating the S minus P arrival times.

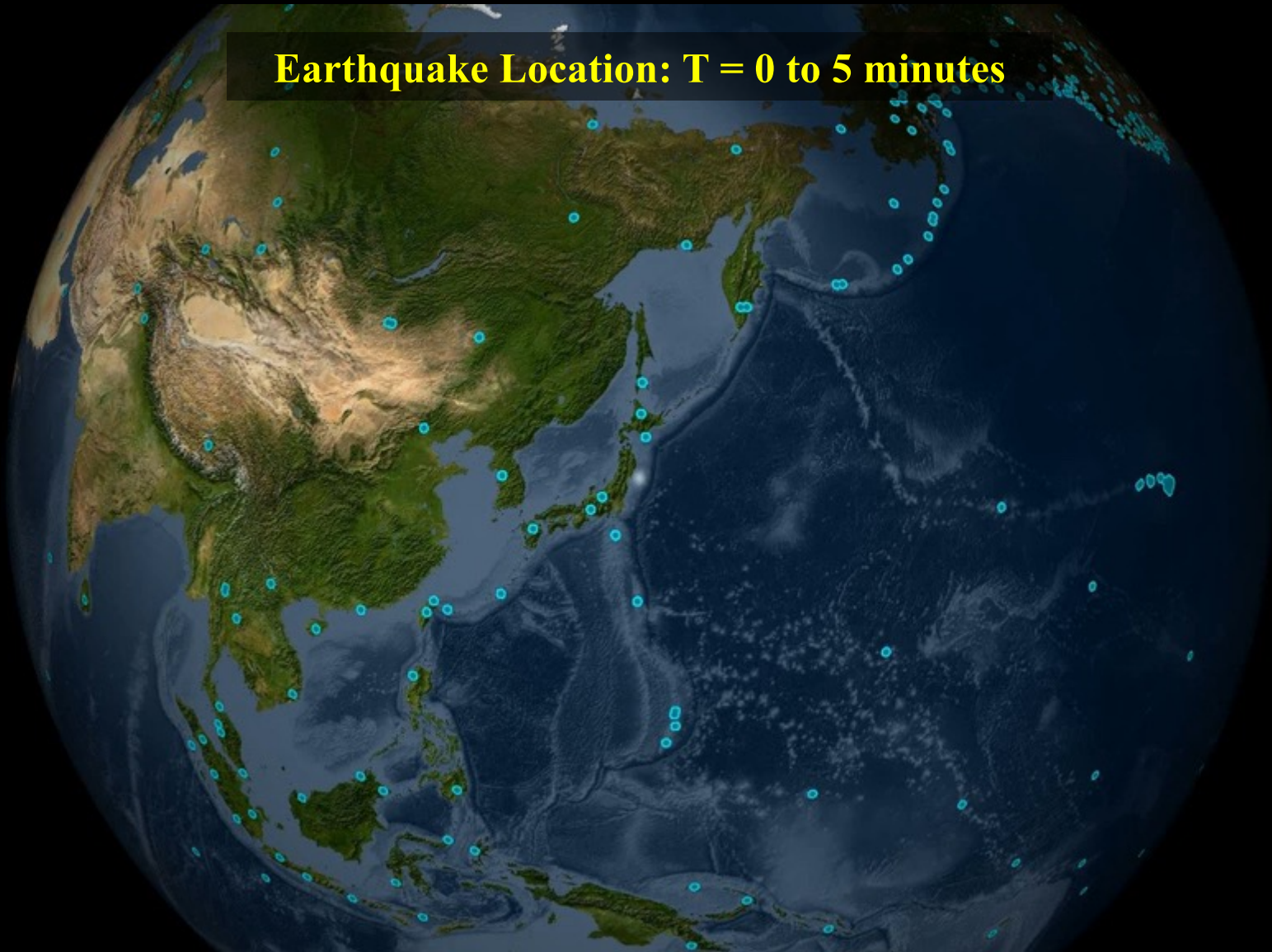
2. Plot on the travel-time graph.

3. Intersection of the circles gives the location.

3.



**Earthquake Location:  $T = 0$  to 5 minutes**



# Earthquake Rupture Complexity

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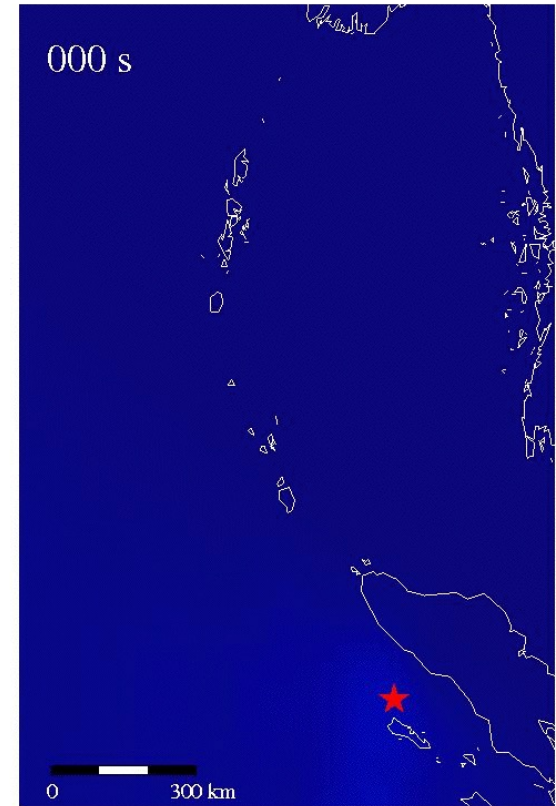
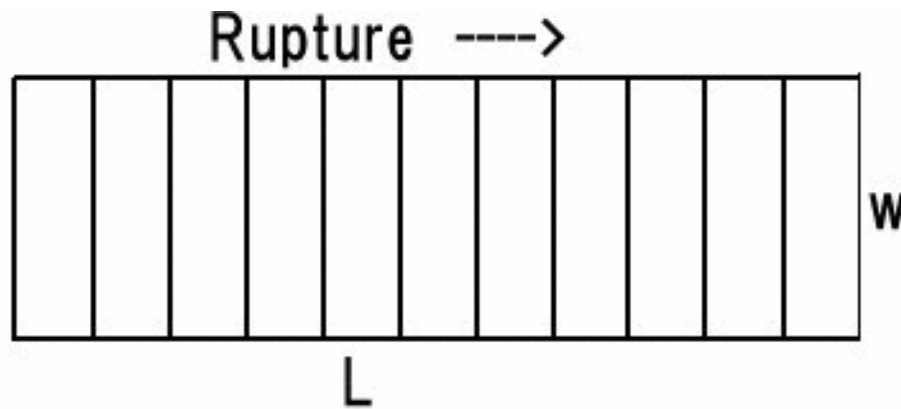
## Great Earthquakes ( $M \geq 8$ )

- Shake for a long time (10s sec to 2-3 minutes)
- Rupture for 100s miles

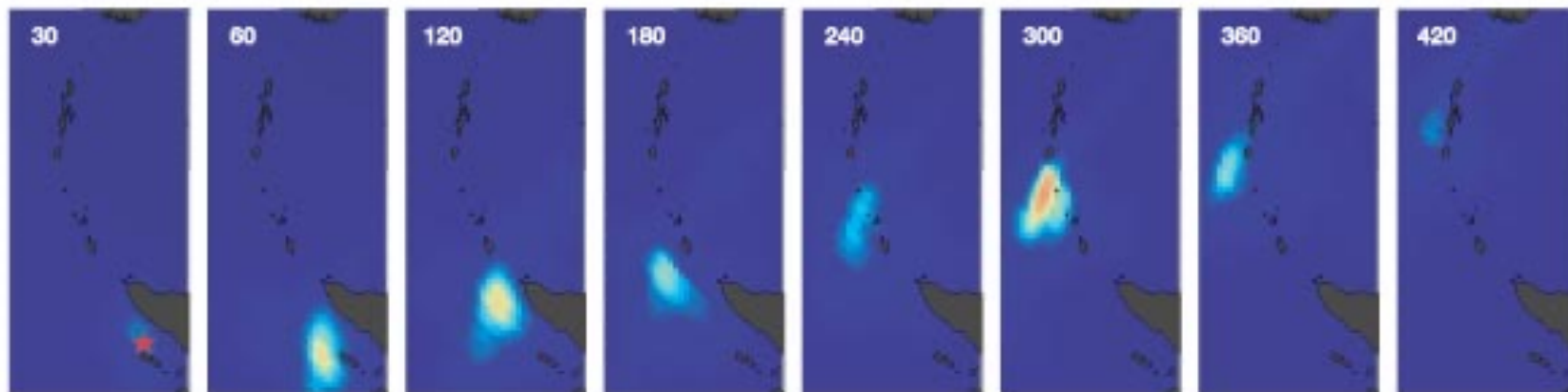


# 2004 Sumatra earthquake

Haskell Line Source  
Dislocation Source



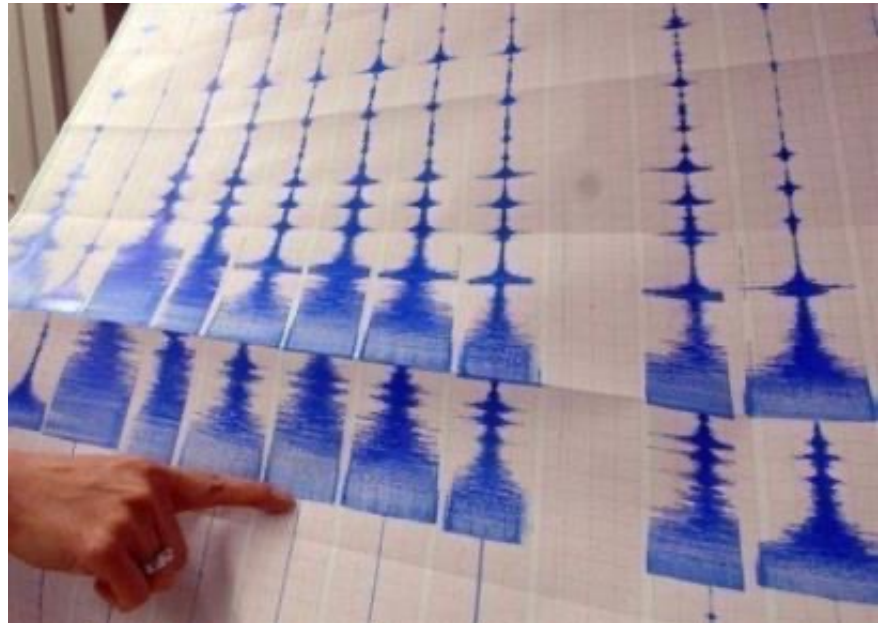
Energy Release imaged by Japan HINET Array



# Earthquake Magnitude & Energy

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M7, Papua New Guinea earthquake







Dispelling myths,  
misconceptions, &  
misunderstandings  
about Earth science



# Earthquake Magnitude VS. Earthquake Intensity

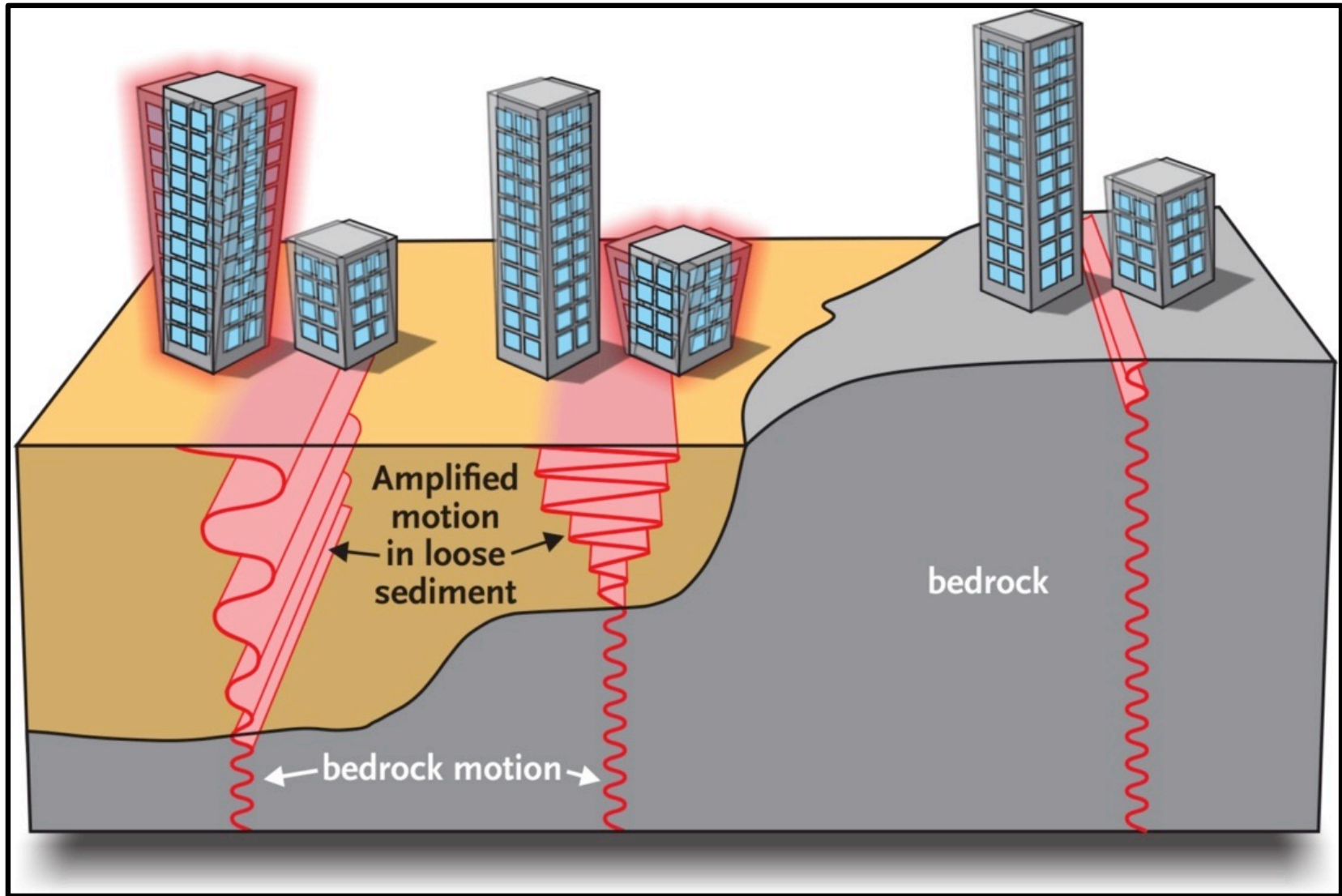
Fixed value vs. variable



[www.iris.edu/earthquake](http://www.iris.edu/earthquake)

INCORPORATED RESEARCH INSTITUTIONS FOR SEISMOLOGY

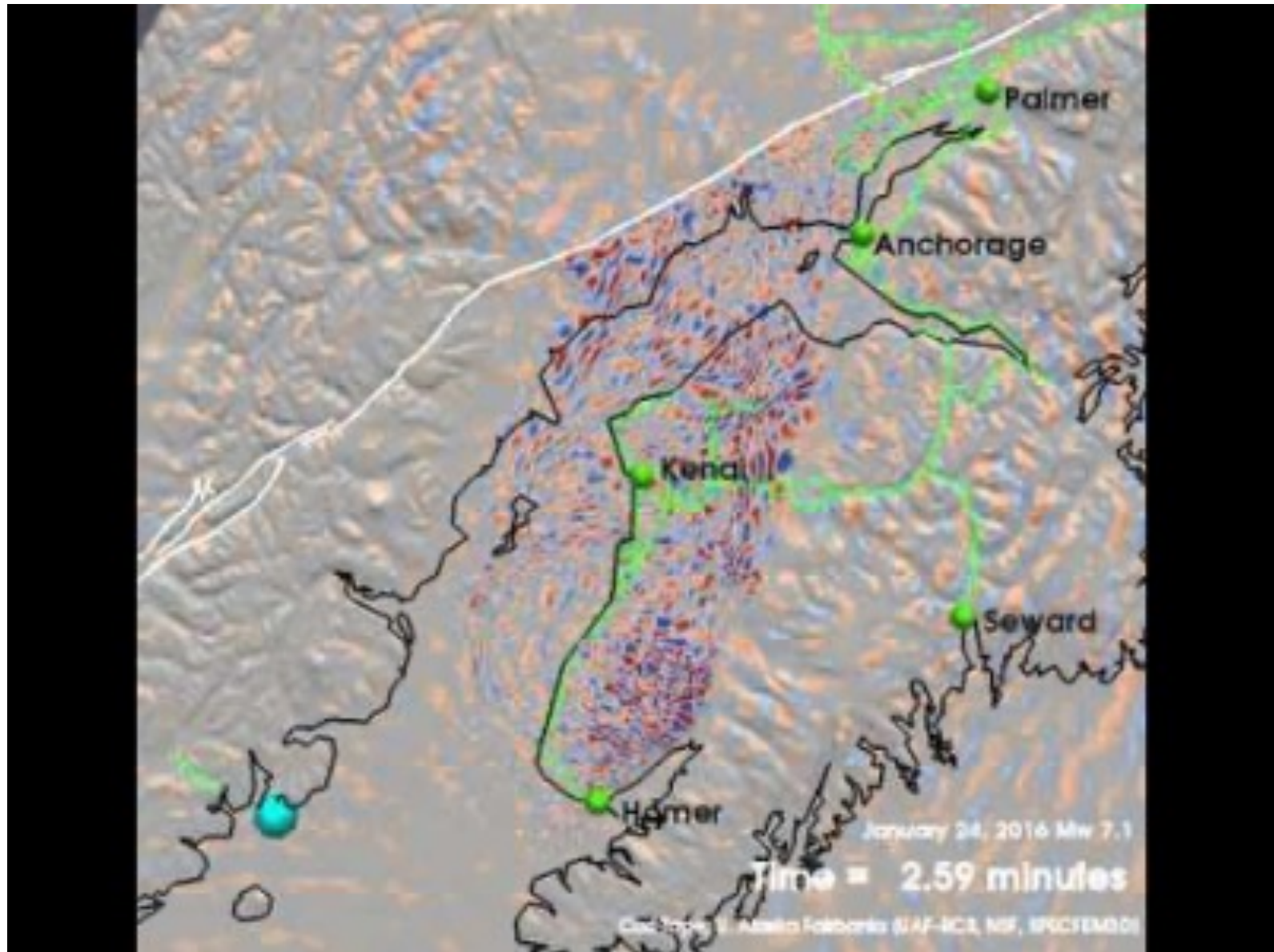
# Ground-Shaking Amplification



Seismic waves are amplified as they pass from bedrock into basins filled with sedimentary rock.

# Site Effects Related to soils and topography

M7.1 Pedro Bay Earthquake (animation by Carl Tape)



# Measuring Earthquake size

- **Historical – Used macroseismic information**
  - Fatalities
  - Maximum shaking
  - Area of intense shaking
- **Did not correlate well from one quake to the next because damage depended on**
  - "True" size (i.e., magnitude)
  - Distance from the epicenter
  - Building design
  - Surface material (rock or dirt) beneath buildings
  - Proximity to populated regions



# Modified Mercalli Intensity Scale

CIIM Intensity	People's Reaction	Furnishings	Built Environment	Natural Environment
I	Not felt			Changes in level and clarity of well water are occasionally associated with great earthquakes at distances beyond which the earthquakes felt by people.
II	Felt by a few.	Delicately suspended objects may swing.		
III	Felt by several; vibration like passing of truck.	Hanging objects may swing appreciably.		
IV	Felt by many; sensation like heavy body striking building.	Dishes rattle.	Walls creak; window rattle.	
V	Felt by nearly all; frightens a few.	Pictures swing out of place; small objects move; a few objects fall from shelves within the community.	A few instances of cracked plaster and cracked windows within the community.	Trees and bushes shaken noticeably.
VI	Frightens many; people move unsteadily.	Many objects fall from shelves.	A few instances of fallen plaster, broken windows, and damaged chimneys within the community.	Some fall of tree limbs and tops, isolated rockfalls and landslides, and isolated liquefaction.
VII	Frightens most; some lose balance.	Heavy furniture overturned.	Damage negligible in buildings of good design and construction, but considerable in some poorly built or badly designed structures; weak chimneys broken at roof line, fall of unbraced parapets.	Tree damage, rockfalls, landslides, and liquefaction are more severe and widespread with increasing intensity.
VIII	Many find it difficult to stand.	Very heavy furniture moves conspicuously.	Damage slight in buildings designed to be earthquake resistant, but severe in some poorly built structures. Widespread fall of chimneys and monuments.	
IX	Some forcibly thrown to the ground.		Damage considerable in some buildings designed to be earthquake resistant; buildings shift off foundations if not bolted to them.	
X			Most ordinary masonry structures collapse; damage moderate to severe in many buildings designed to be earthquake resistant.	

# USGS ShakeMap 2010 M<sub>w</sub> 8.8 Maule, Chile

Macroseismic Intensity Map USGS  
ShakeMap: offshore Bio-Bio, Chile  
Feb 27, 2010 06:34:11 UTC M8.8 S36.12 W72.90 Depth: 22.9km  
ID:official20100227063411530\_30

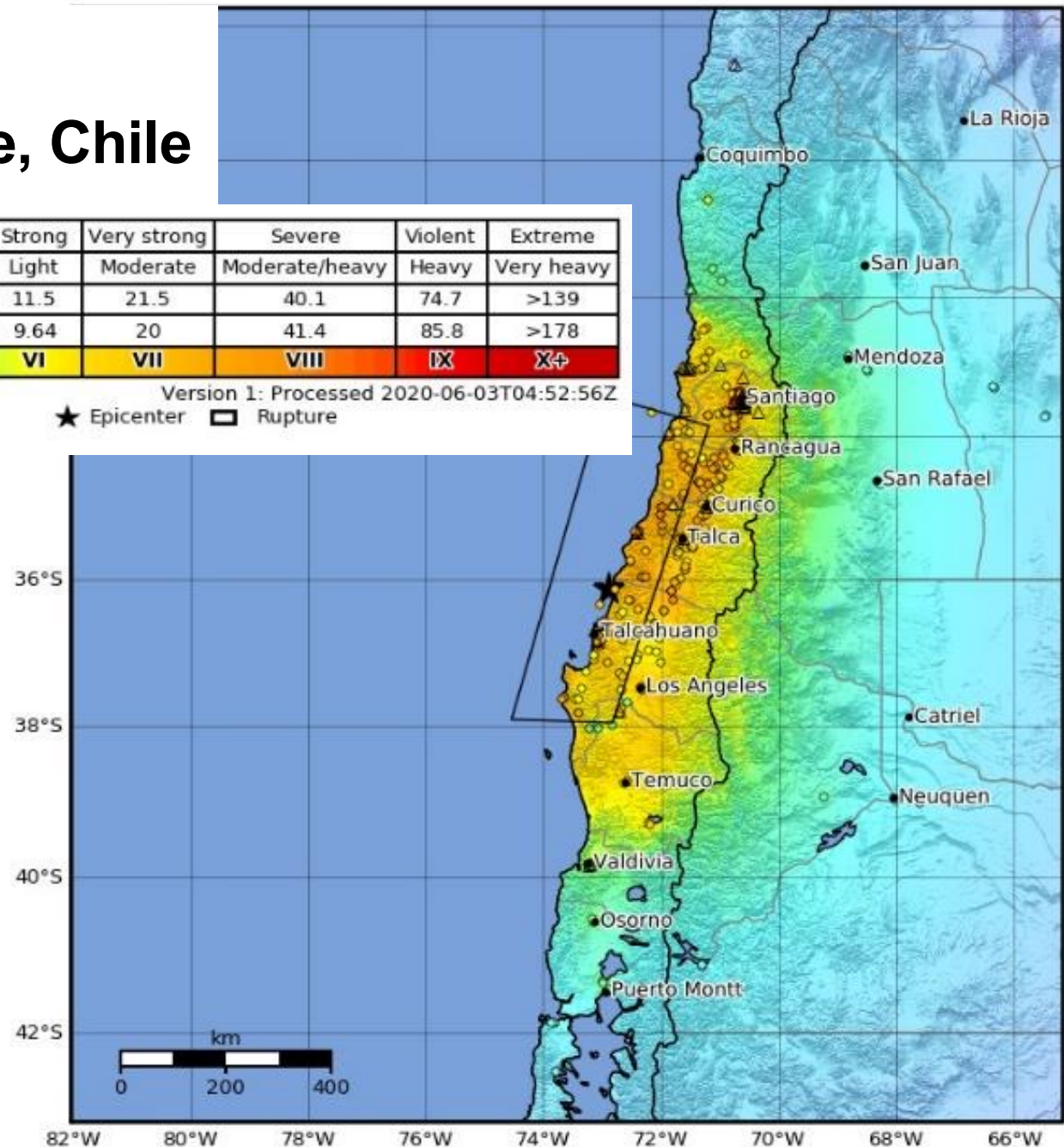
SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	None	None	None	Very light	Light	Moderate	Moderate/heavy	Heavy	Very heavy
PGA(%g)	<0.0464	0.297	2.76	6.2	11.5	21.5	40.1	74.7	>139
PGV(cm/s)	<0.0215	0.135	1.41	4.65	9.64	20	41.4	85.8	>178
INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+

Scale based on Worden et al. (2012)

△ Seismic Instrument ○ Reported Intensity

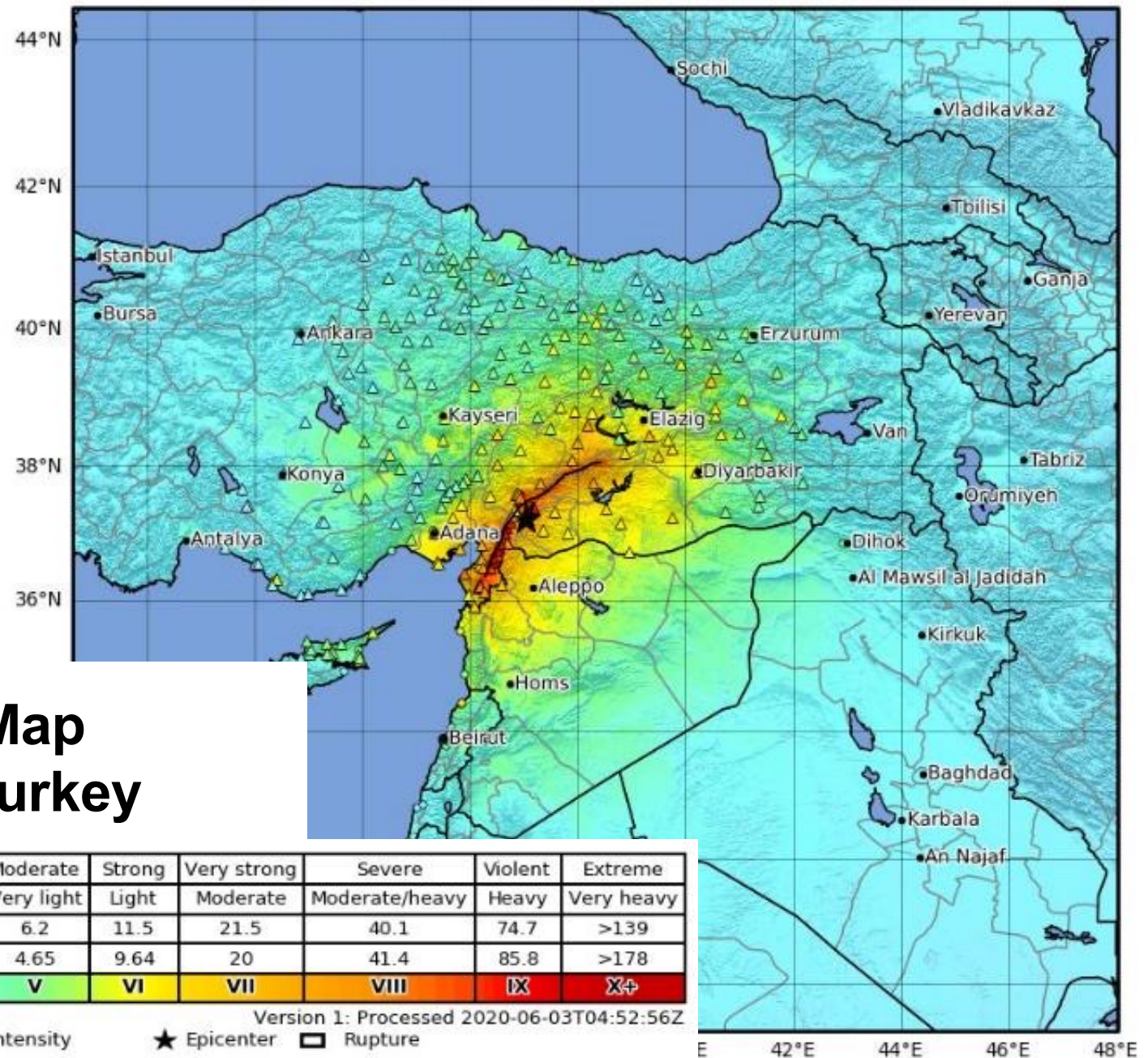
★ Epicenter □ Rupture

Version 1: Processed 2020-06-03T04:52:56Z





Macroseismic Intensity Map USGS  
 ShakeMap: 25 km ENE of Nurdağı, Gaziantep, TR  
 Feb 06, 2023 01:17:34 UTC M7.8 N37.23 E37.01 Depth: 10.0km ID:us6000jllz



## USGS ShakeMap 2023 M<sub>w</sub> 7.8 Turkey

SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	None	None	None	Very light	Light	Moderate	Moderate/heavy	Heavy	Very heavy
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Version 1: Processed 2020-06-03T04:52:56Z

# Earthquake Magnitude

General form:

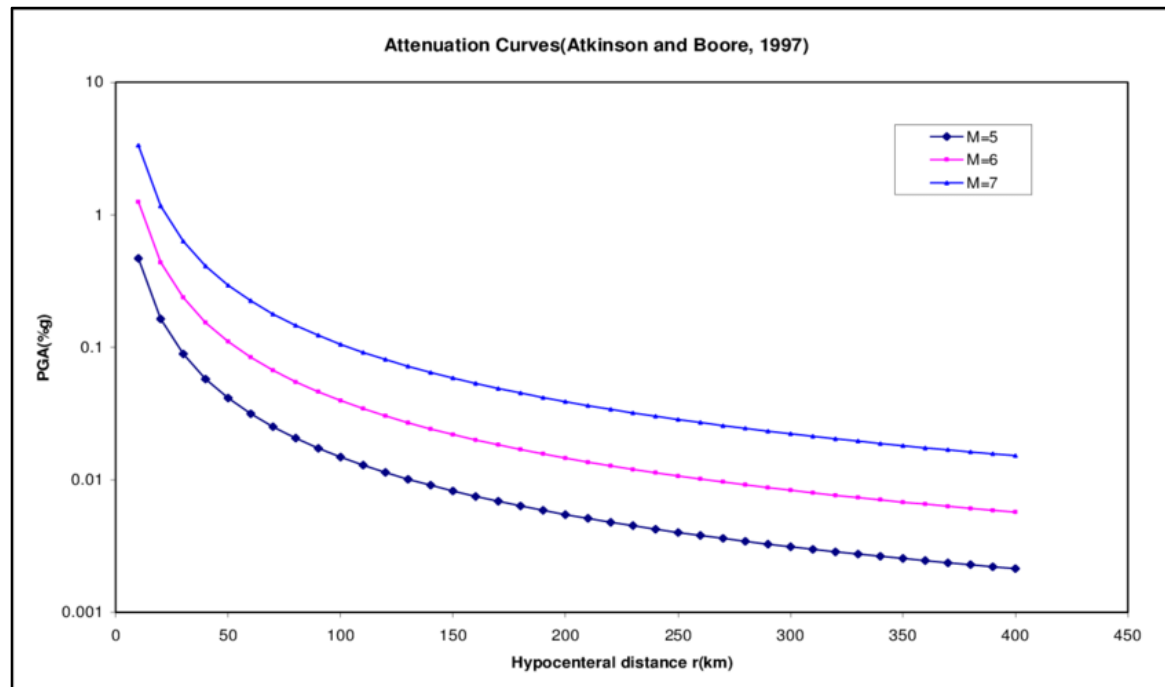
$$M = \log(A/T) + F(h, \Delta) + C$$

Amplitude

Period

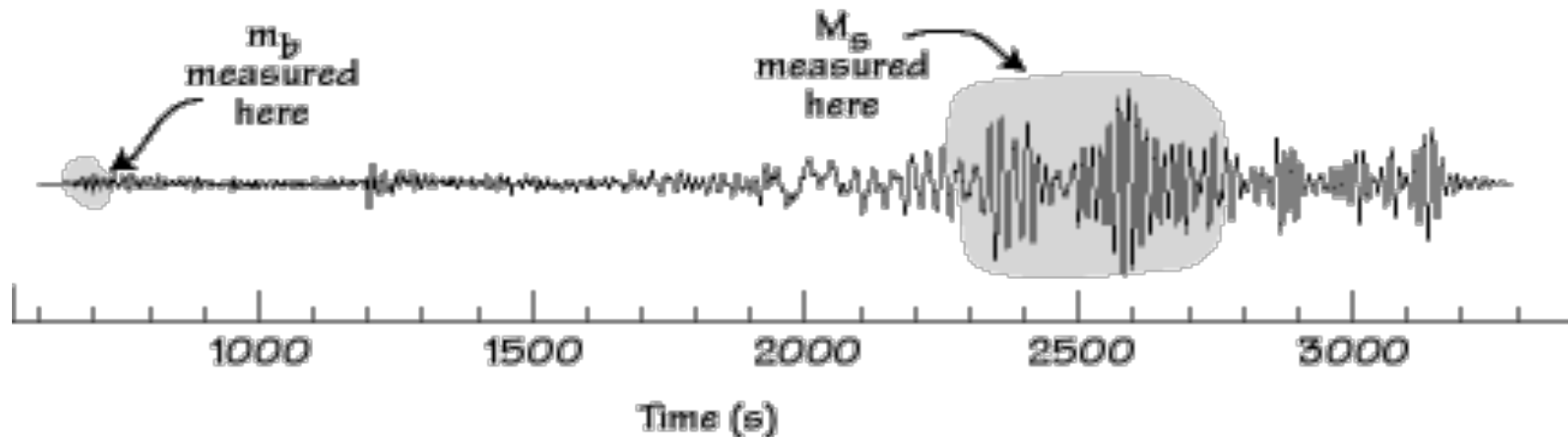
Correction for depth  
and distance

Regional scaling factor



# Richter and Gutenberg's Teleseismic (distant) magnitudes

## Body wave Magnitude $m_b$ and Surface wave and $M_s$



$$m_b = \log (A/T) + Q(D,h)$$

**T:** period (secs),  $0.1 \leq T \leq 3.0$

**A:** P wave amplitude (microns)  
(not necessarily the maximum)

**Q:** scale factor ( $D \geq 5^\circ$ )

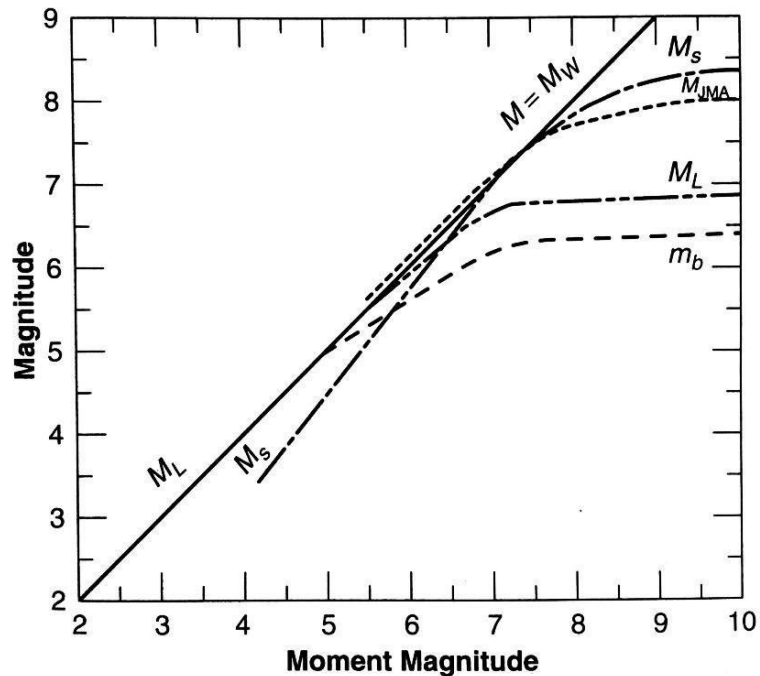
$$M_s = \log (A/T) + 1.66 \log D + 3.3$$

**A:** maximum amplitude (microns)  
vertical component of the  
surface wave within the period  
range  $18 \leq T \leq 22$ .

**D:**  $20^\circ \leq D \leq 160^\circ$ .

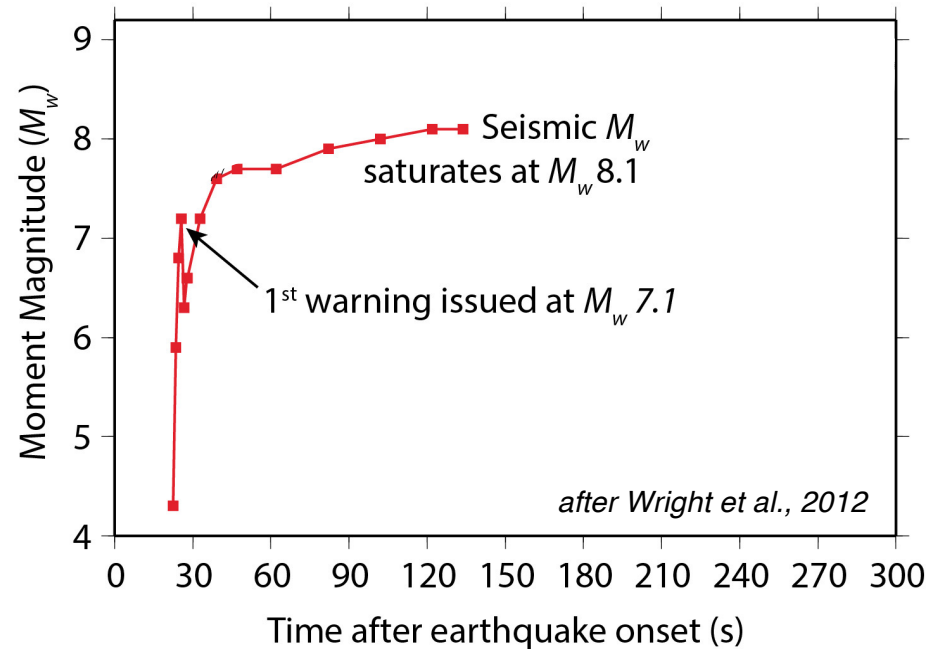
No depth corrections!

# Saturation!



*Idriss, 1985*

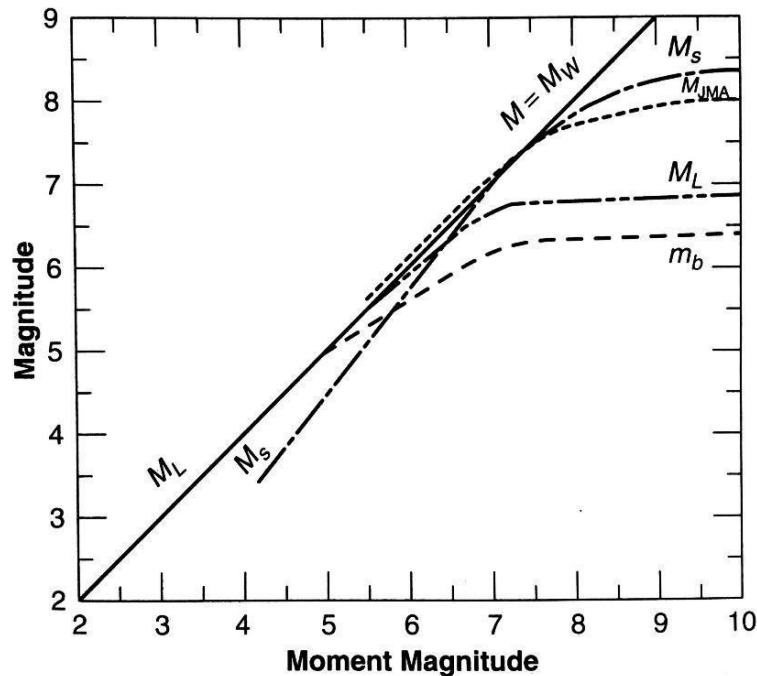
2011 Tohoku-Oki (Japan) Earthquake



## When does this happen?

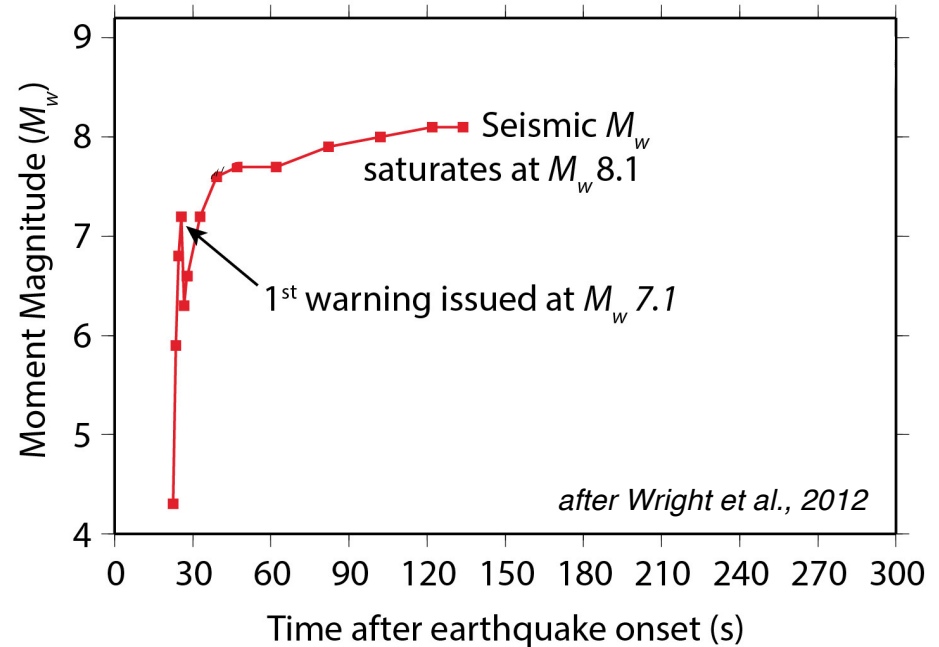
- 1) Time window saturation – when time window is less than duration of rupture (particularly affects  $m_b$ )
- 2) Spectral saturation – when wavelength too short to see entire rupture (affects  $m_b$ ,  $M_L$ , and  $M_s$ )

# Saturation!



Idriss, 1985

2011 Tohoku-Oki (Japan) Earthquake



## How do we overcome this problem?

- Examine longer period waves
- $M_w$ ,  $M_{wp}$ , Mantle magnitude ( $M_m$ ),  
Centroid Moment Tensor (CMT) – moment magnitude
- GNSS/GPS data (next talk)

# Types of Magnitude Scales

			Period Range
$M_L$	Local magnitude (California)	regional S & surface waves	0.1-1 sec
$M_j$	JMA (Japan Meteorol. Agency)	regional S & surface waves	5-10 sec
$m_b$	Body wave magnitude	teleseismic P waves	1-5 sec
$M_s$	Surface wave magnitude	teleseismic surface waves	20 sec

The methods below overcome the effects of saturation:

$M_{wp}$	P-wave moment magnitude	teleseismic P waves	10-60 sec
$M_w$	Moment magnitude	teleseismic surface waves	> 200 sec
$M_m$	Mantle magnitude	teleseismic surface waves	> 200 sec

*Ref: USGS Seismology and Tsunami Warnings, 2006 (Earthquake Source)*



# Moment magnitude (Mw)

$$M_w = (2/3) \log_{10} (M_0) - 16.1$$

- Introduced in 1979 by Hanks and Kanamori
- Based on source parameter  $M_0$  and is not frequency dependent, **does NOT saturate**
- Based on earthquake **energy release**
- **Related to fault slip** and not ground shaking
- Used to estimate the magnitude of **large earthquakes**
- Very useful for **tsunami modeling**

# Moment Magnitude

Understanding moment magnitude  
...or what bumped the Richter scale?



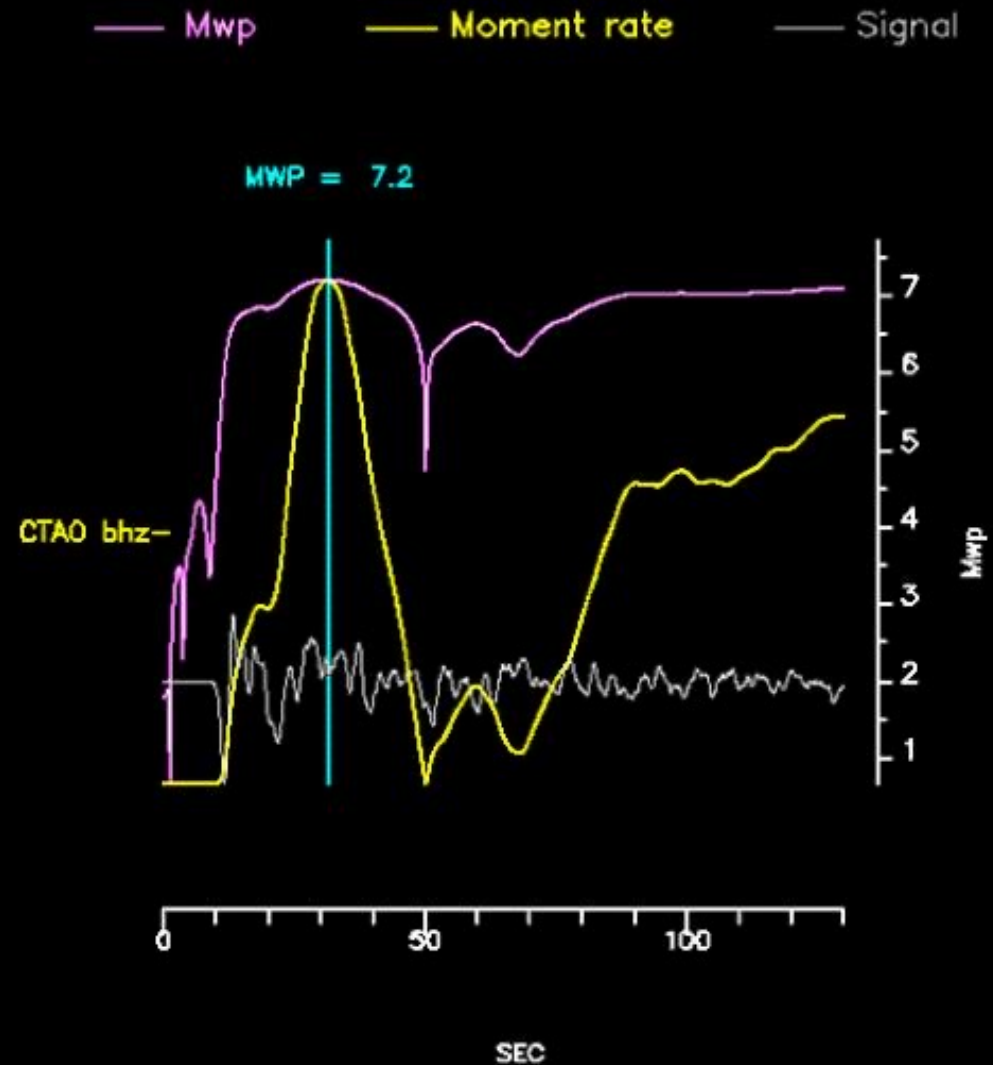
IRIS

# Mwp Magnitude

- Moment magnitude based on **initial long-period P-waves**
- Developed by S. Tsuboi and others (1995)
- **Empirical estimate of the moment magnitude**
  - integrate the vertical velocity from a seismogram
- **Accurate results within 3-4 minutes** of OT (P, pP arrivals)
- Primary initial **magnitude estimate at PTWC for  $M > 6$**  (replaced  $M_s$ )
- **Subject to site and path effects, source complexity, contamination from other large earthquakes**

**$M_{wp}$**  [Tsuboi et al., 1995; 1999]

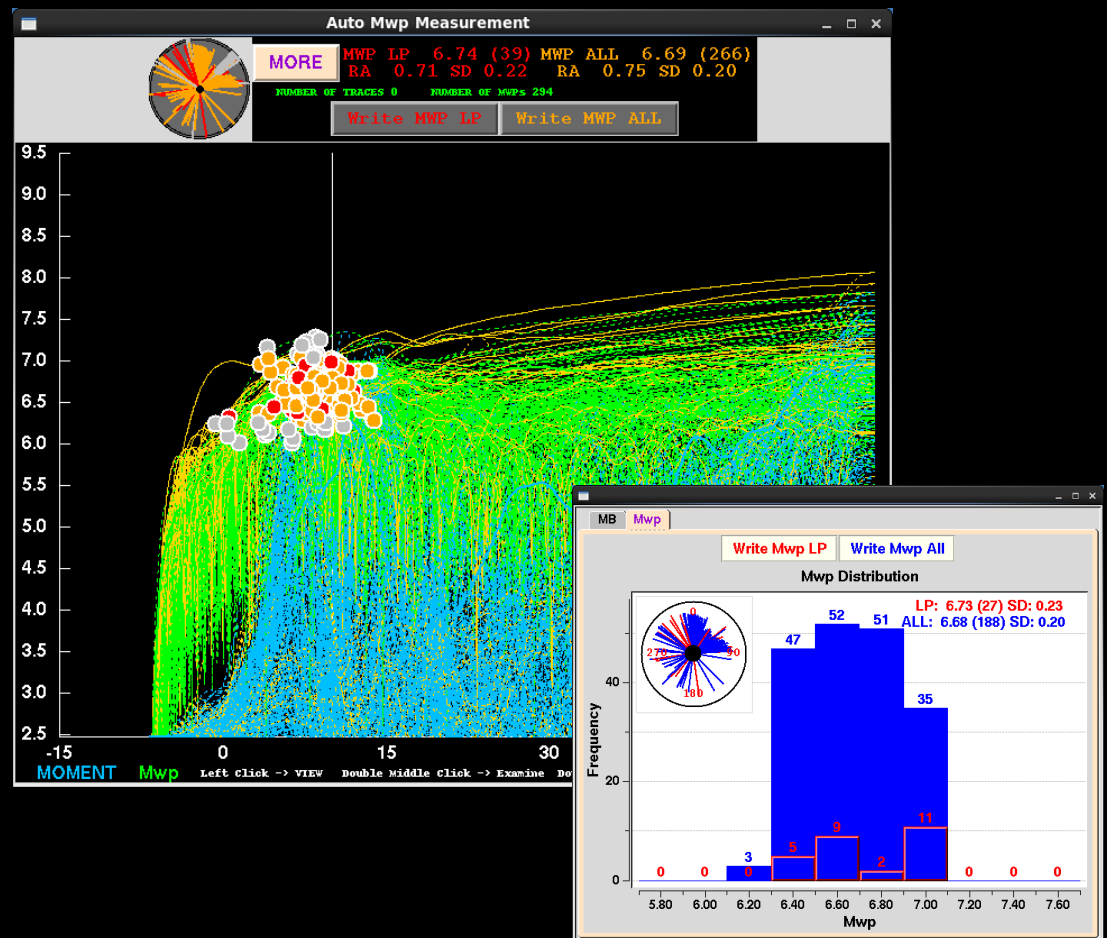
- Double integration of  $v(t)$   
 $v(t) \rightarrow Mo(t) \rightarrow Mw(t)$
- Peak  $Mw(t) \rightarrow Mwp$
- Fast; less prone to **saturation**





**$M_{wp}$**  [Tsuboi et al., 1995; 1999]

- Double integration of  $v(t)$   
 $v(t) \rightarrow Mo(t) \rightarrow Mw(t)$
- Peak  $Mw(t) \rightarrow Mwp$
- Fast; less prone to saturation



# TWC magnitude determination – PTWC (5 primary)

- **ML** - local magnitude, HI/PR/USVI ( $2 < M < 6$ )
- **mb** - body wave magnitude, largest P-waves ( $4 < M < 6.5$ )
- **Mm** - mantle wave magnitude, 50-400s surface waves ( $M > 6$ )
- **Mwp** – estimate of moment magnitude from integrated P-waves ( $5 < M < 8$ )
- **W-phase** - long-period phase, gives stable results, does not saturate, but takes ~20-25 minutes ( $M > 5$ )

# TWC magnitude determination – PTWC (5 primary)

- Measurements are not precise.  
Different magnitudes result from using different scales/methods and datasets.
- **Many methods saturate for large earthquakes**
- **$M_0 \rightarrow$  EQ energy  $\rightarrow$  "true" EQ size but takes time**
- **TWC's rely on  $M_{wp}$  for preliminary message products, method can still underestimate great earthquake sizes.**
- **WCMT provides fault geometry and authoritative estimate of  $M_w$  but takes time (~15-25 minutes)**



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7-18 August 2023, Honolulu, Hawaii USA

# Thank You

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